



Empirical-Statistical Downscaling & Record-Statistics



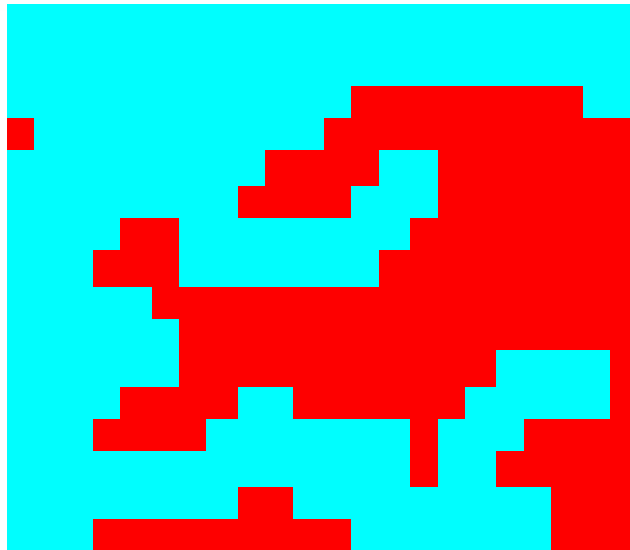
R.E. Benestad

Outline

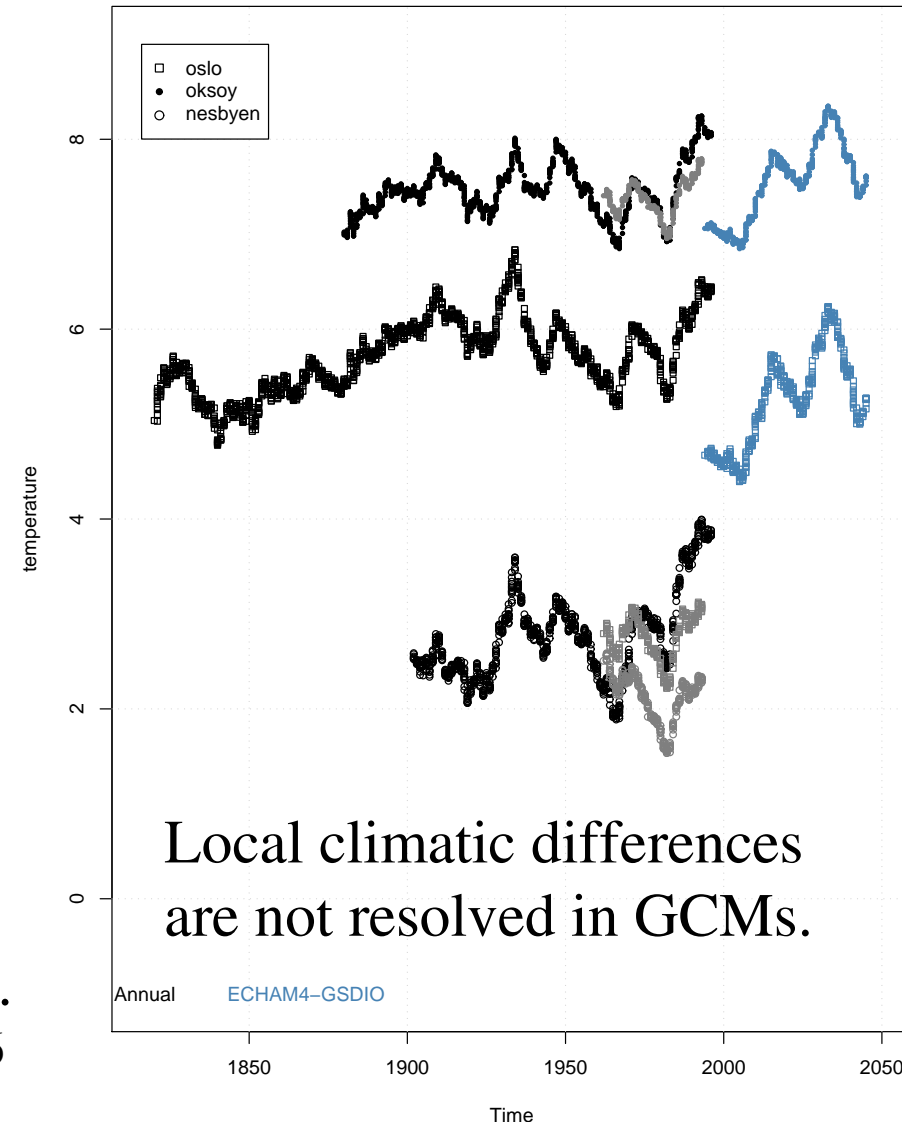


- Theoretical basis.
 - Why do we need downscaling?
- Methods.
 - Dynamical v.s. empirical-statistical
 - The common EOF frame-work
 - Statistical models: Regression, CCA, SVD, ...
 - Empirical-statistical: e.g linear, analog
- Uncertainties.
- Extremes & record values.
 - Downscaling pdfs
 - Exponential law & dependency on local conditions
 - Record-events: test for non-stationarity
 - 'Extremes' as in 'severe' events

Why downscaling?



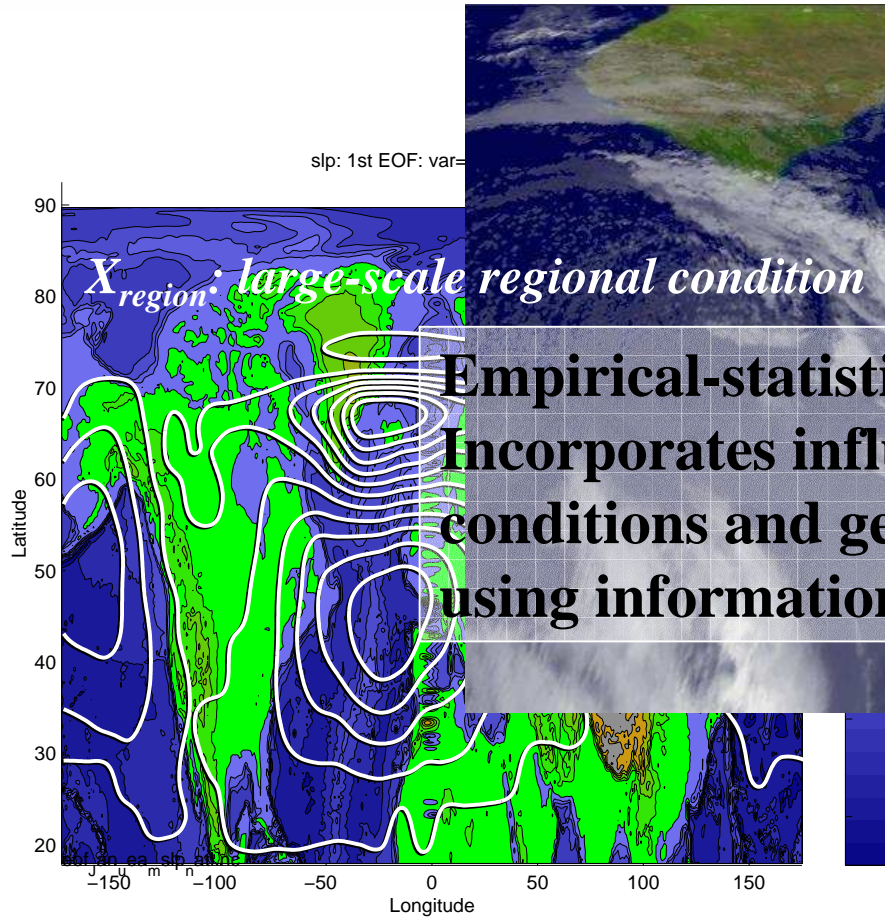
Interpolated Temperatures v.s. station Observations



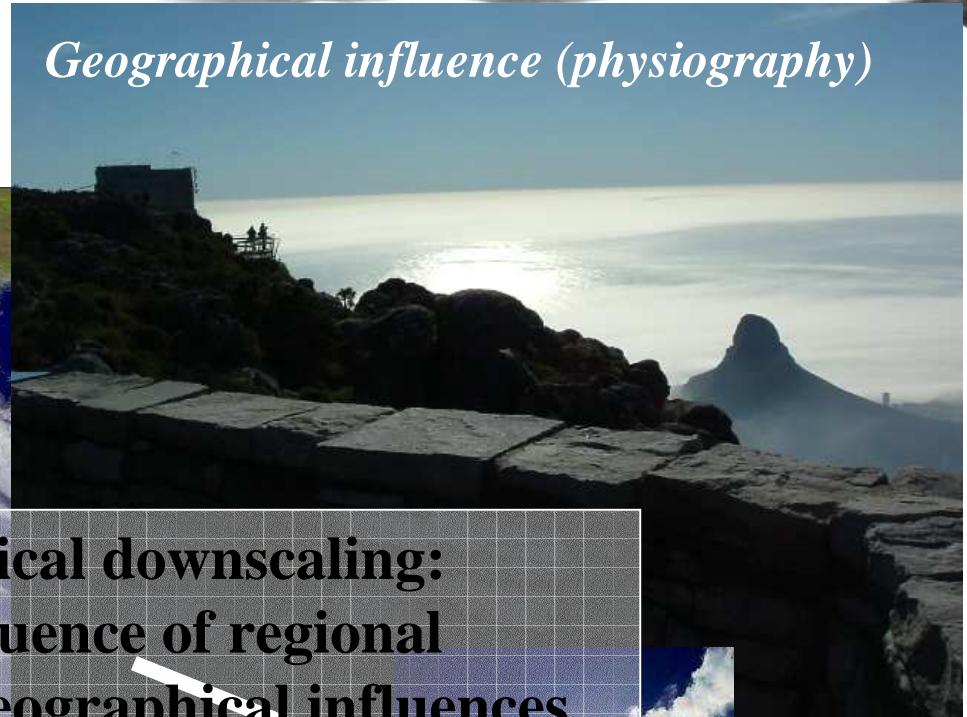
“Skillful spatial scale”: ~ 8 grid-pts.
Grotch & McCracken (1991), *J. Clim.*, **4**, p. 286

Principles of Downscaling

Large-scale (GCMs, re-analysis)

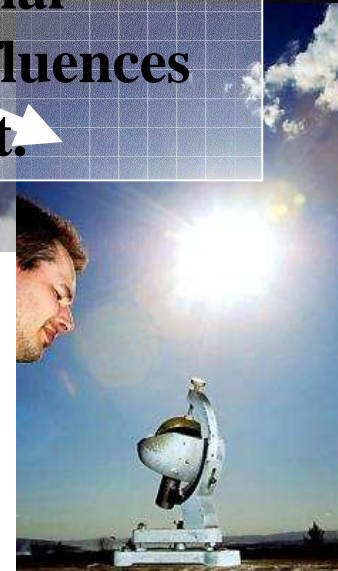


Geographical influence (physiography)



**Empirical-statistical downscaling:
Incorporates influence of regional
conditions and geographical influences
using information from the past.**

$$X_{local} = \psi(X_{region}, \text{physiography})$$

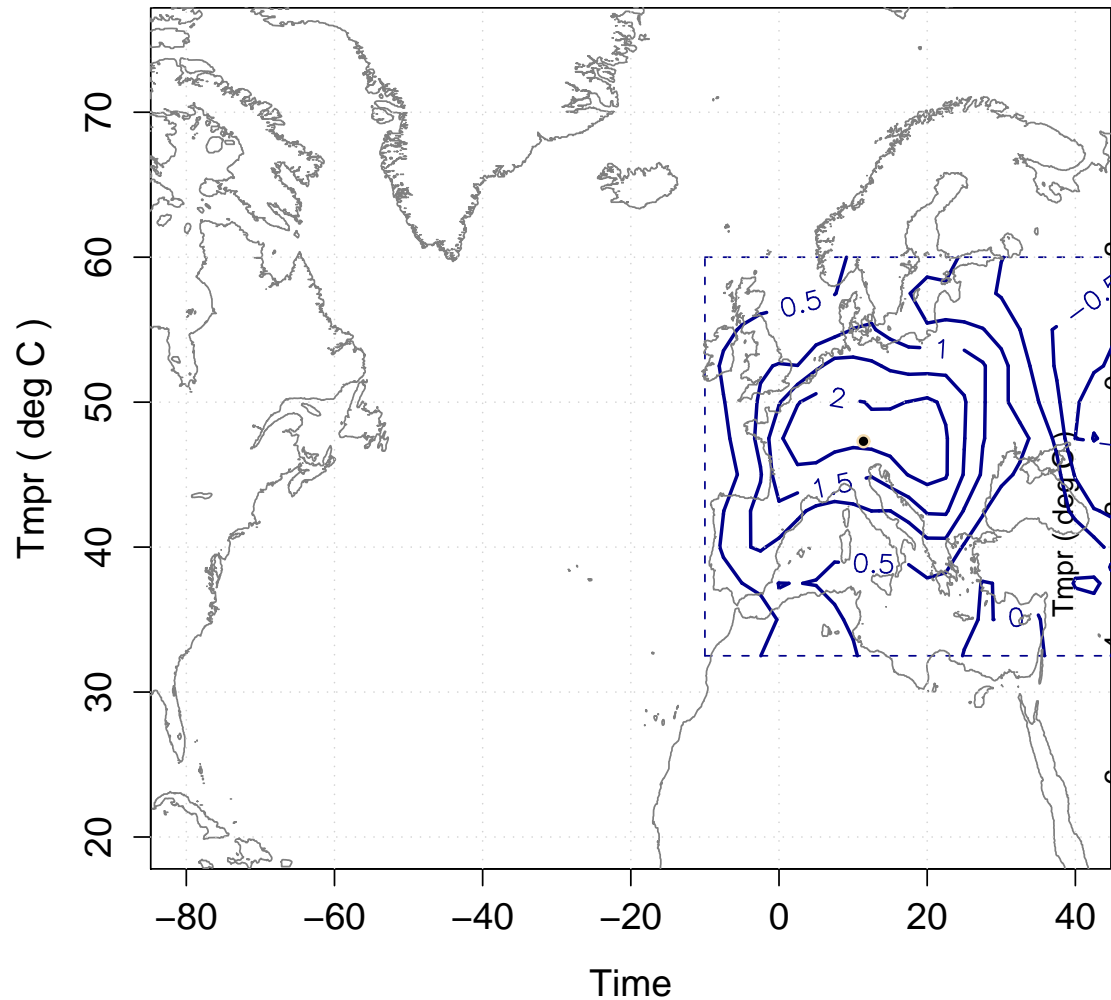


Small-scale
(Direct measurements)

Regression example for Innsbruck...

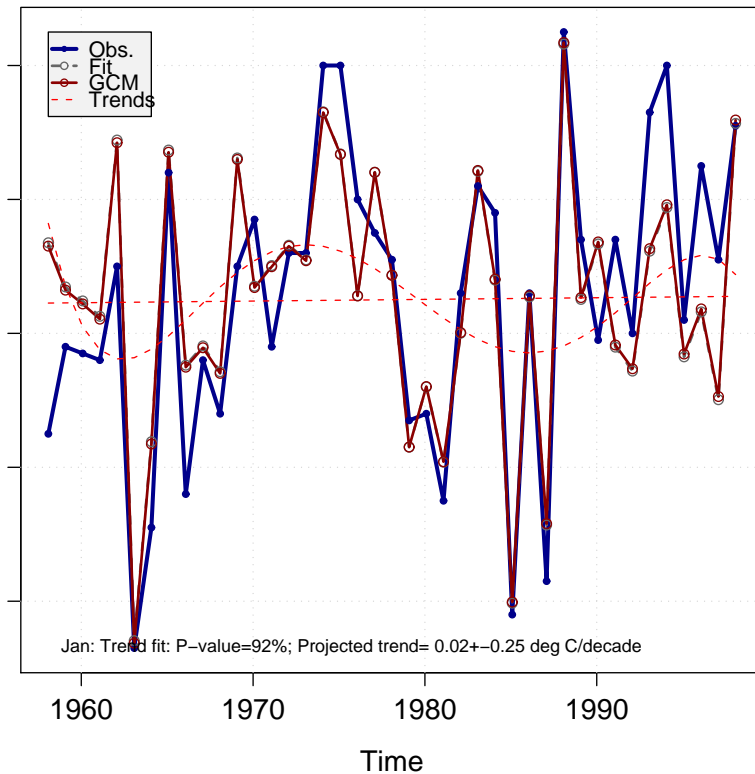


Empirical Downscaling (ncep_t2m [10W50E-30N60N] -> Tmpr)

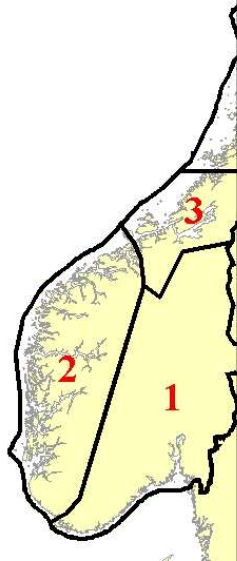


Calibration: Jan Tmpr at INNSBRUCK using ncep_t2m: R2=81%, p-value=0%.

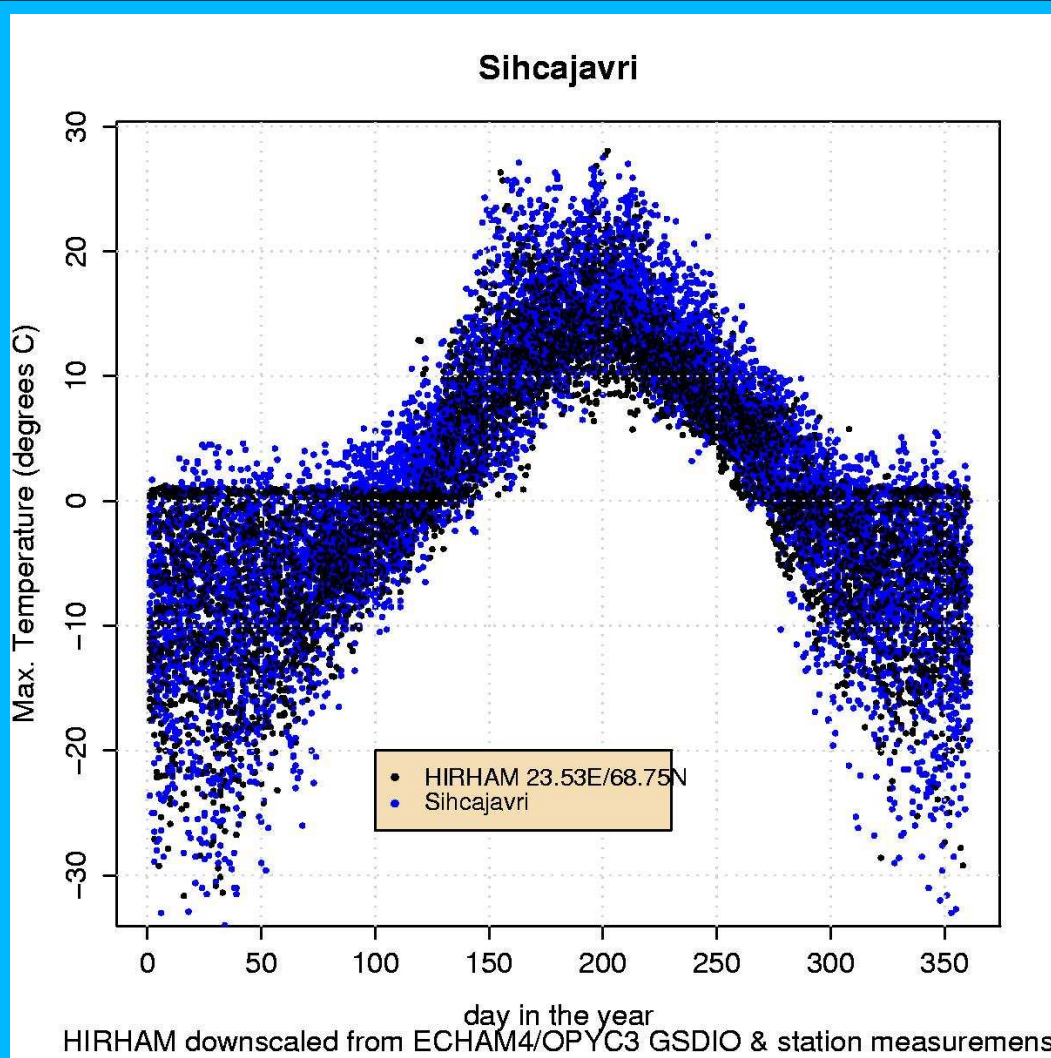
Empirical Downscaling (ncep_t2m [10W50E-30N60N] -> Tmpr)



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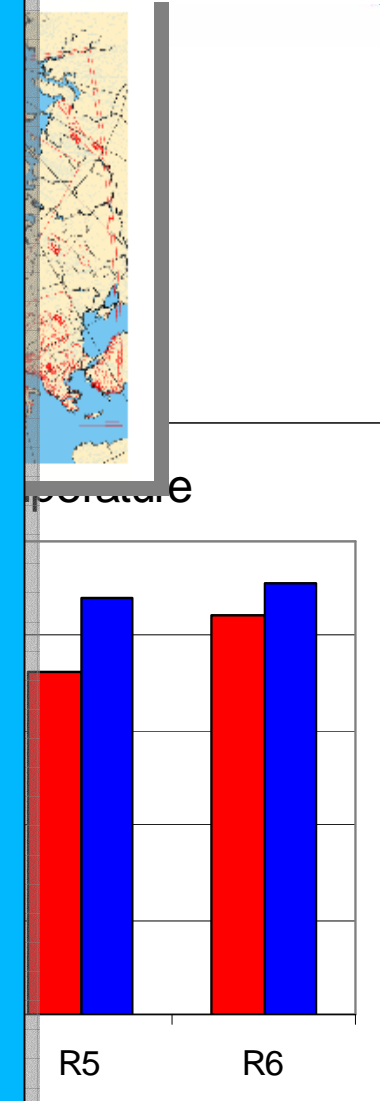


Hanssen-I



Dynamical DS not necessarily 'better' than empirical-statistical. Stationarity-problems associated with parameterisation (statistical) and not more 'physically consistent' (systematic biases – also see figure!).

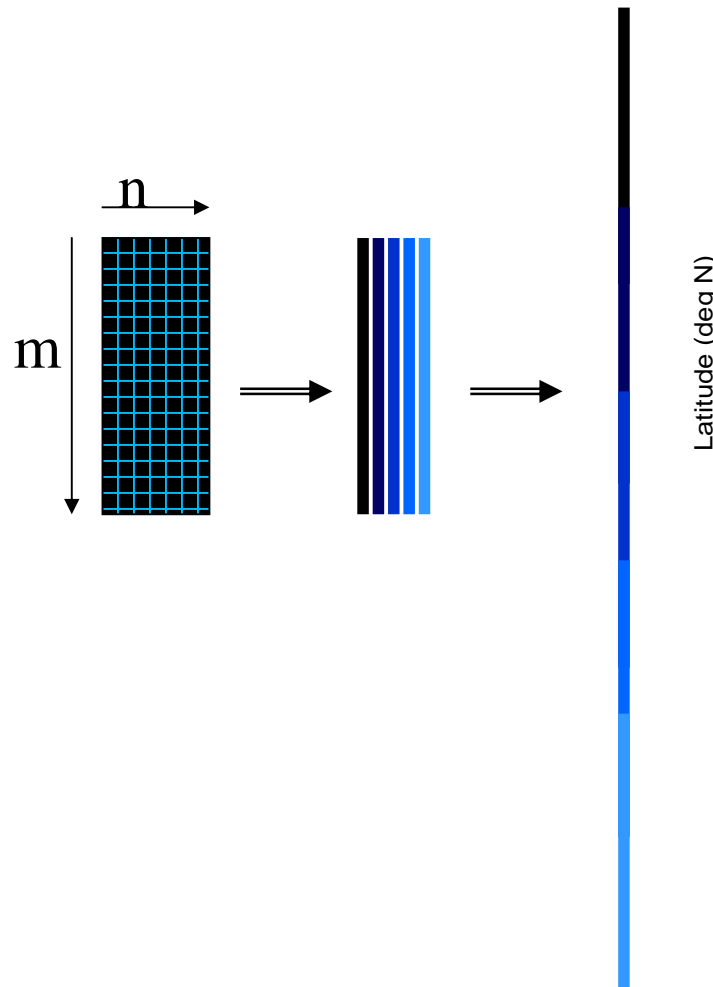
ling: 2 completely
g strategies.





2-dimensional data matrix converted to a 1D vector:

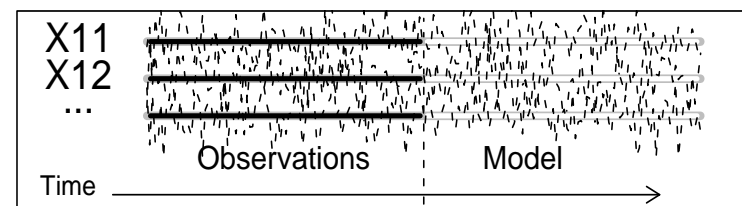
$$Y_{ij} \rightarrow \tilde{y} \in \mathcal{R}^{n \times m}$$



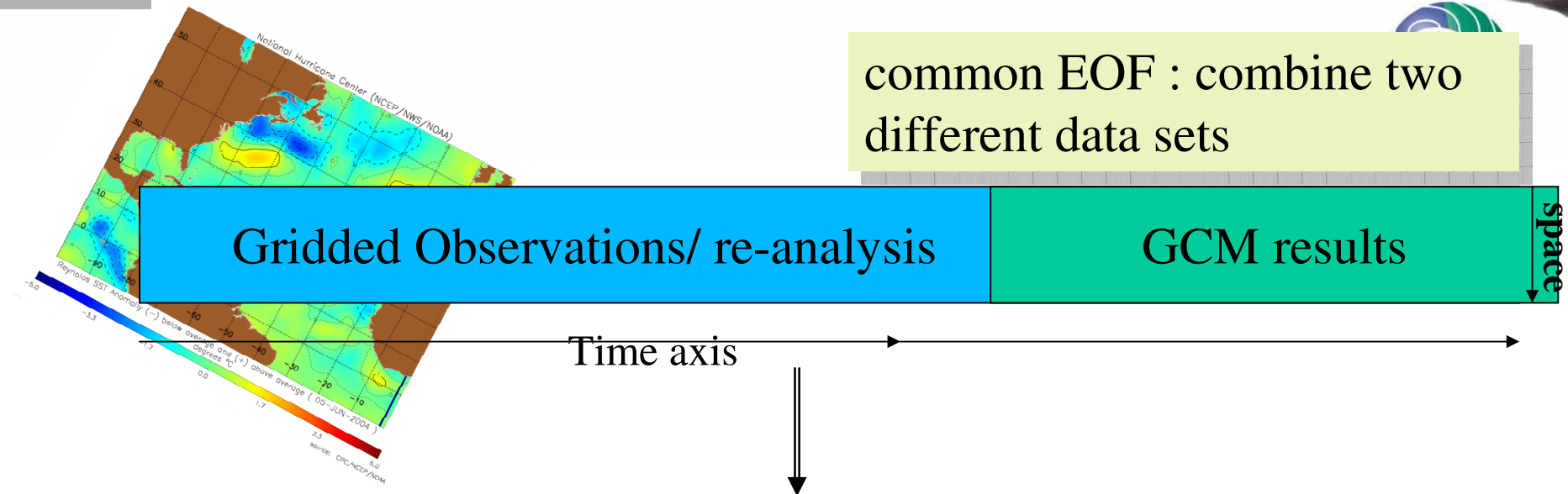
Example of data grid

X14	X24	X34	X44	X54	X64	X74	X84	X94
X13	X23	X33	X43	X53	X63	X73	X83	X93
X12	X22	X32	X42	X52	X62	X72	X82	X92
X11	X21	X31	X41	X51	X61	X71	X81	X91

Longitude (deg E)

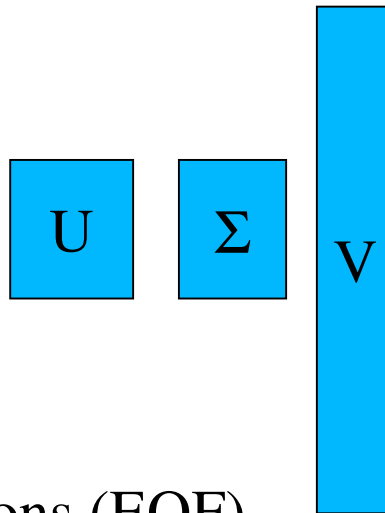


A question of how to organize the data...



PCA: Singular Vector Decomposition (SVD):

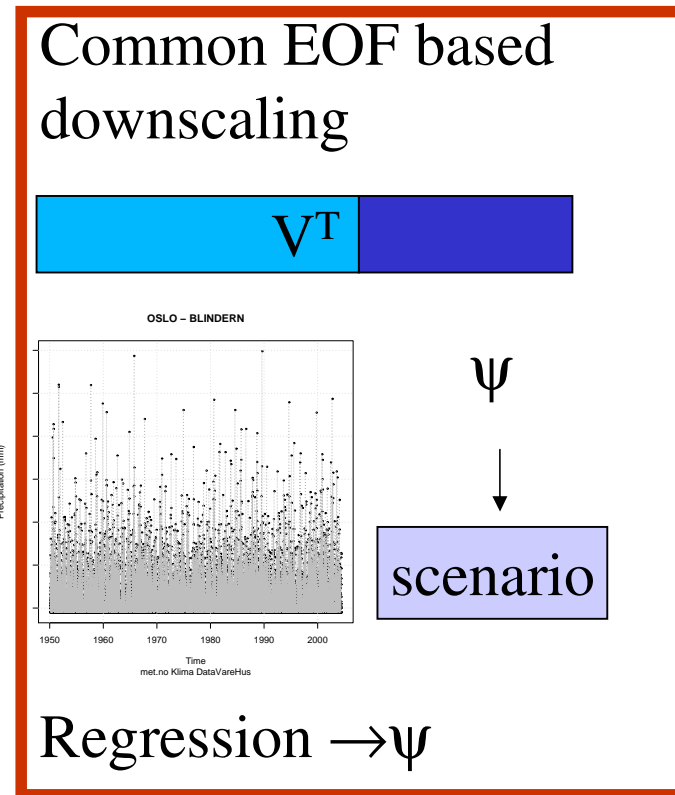
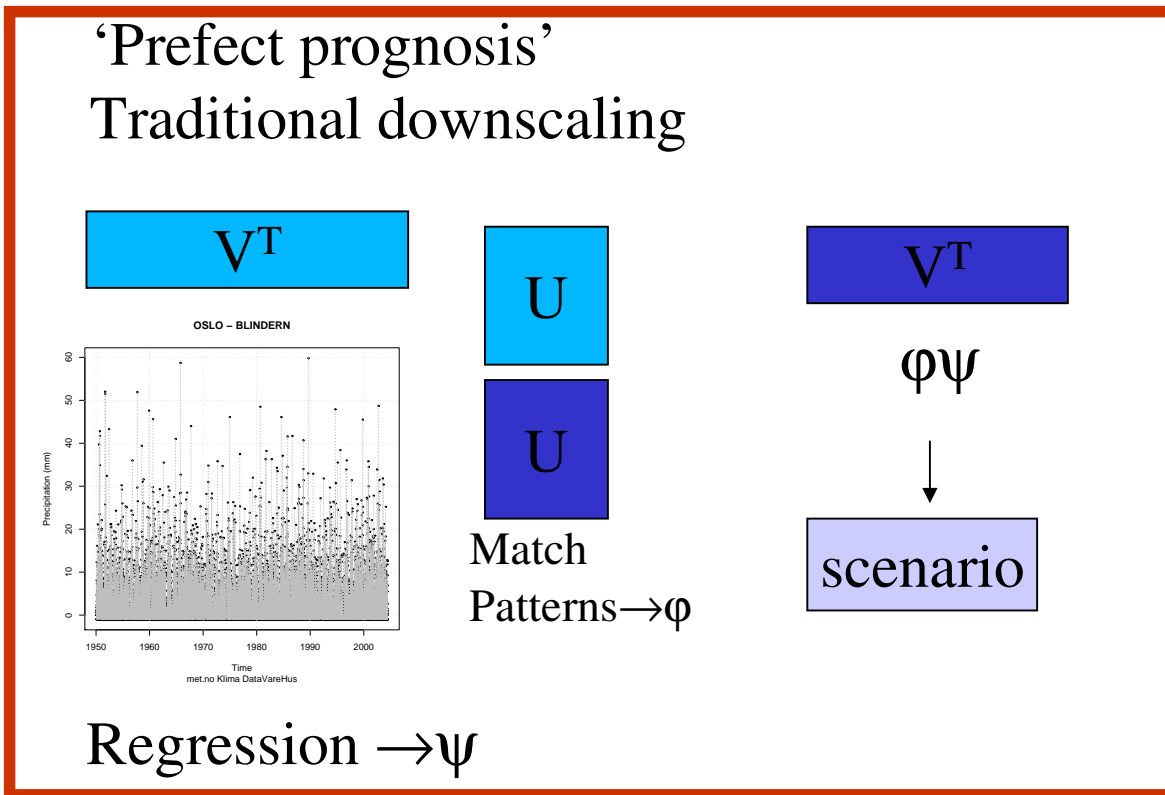
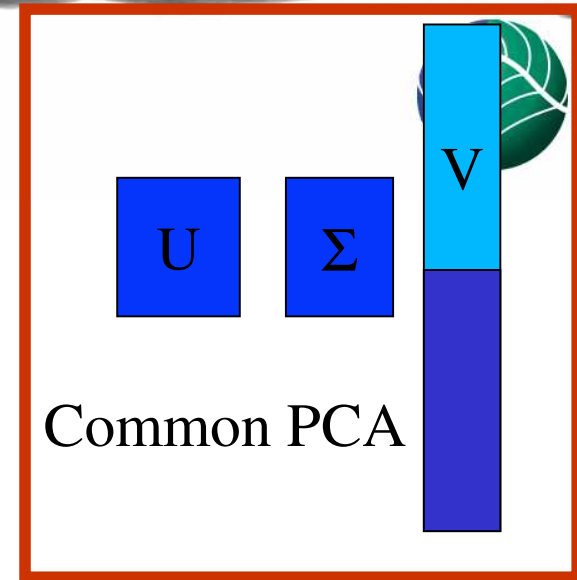
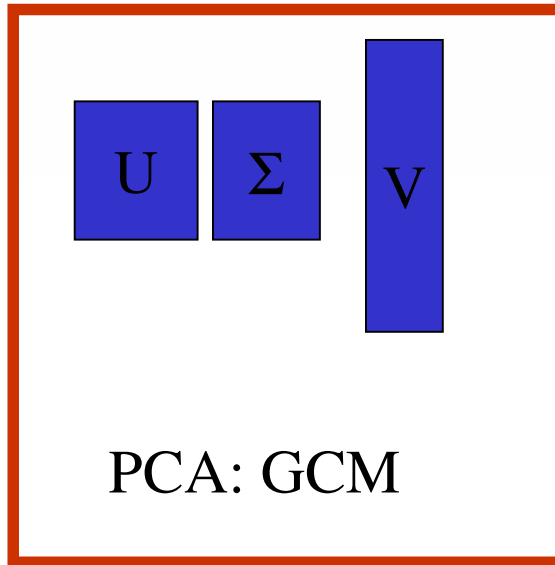
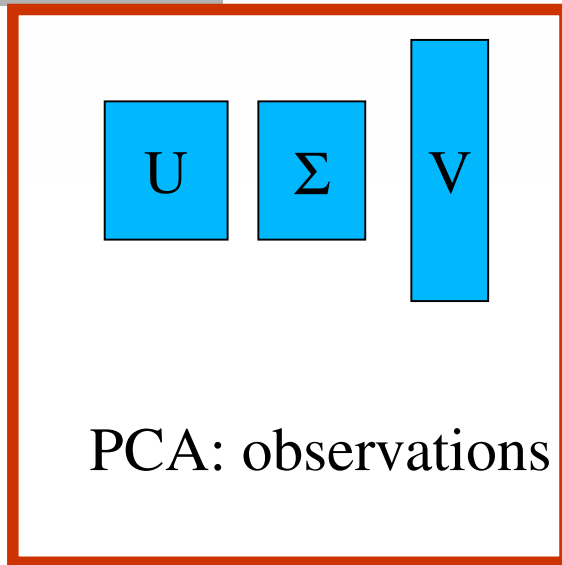
$$X = U \Sigma V^T$$



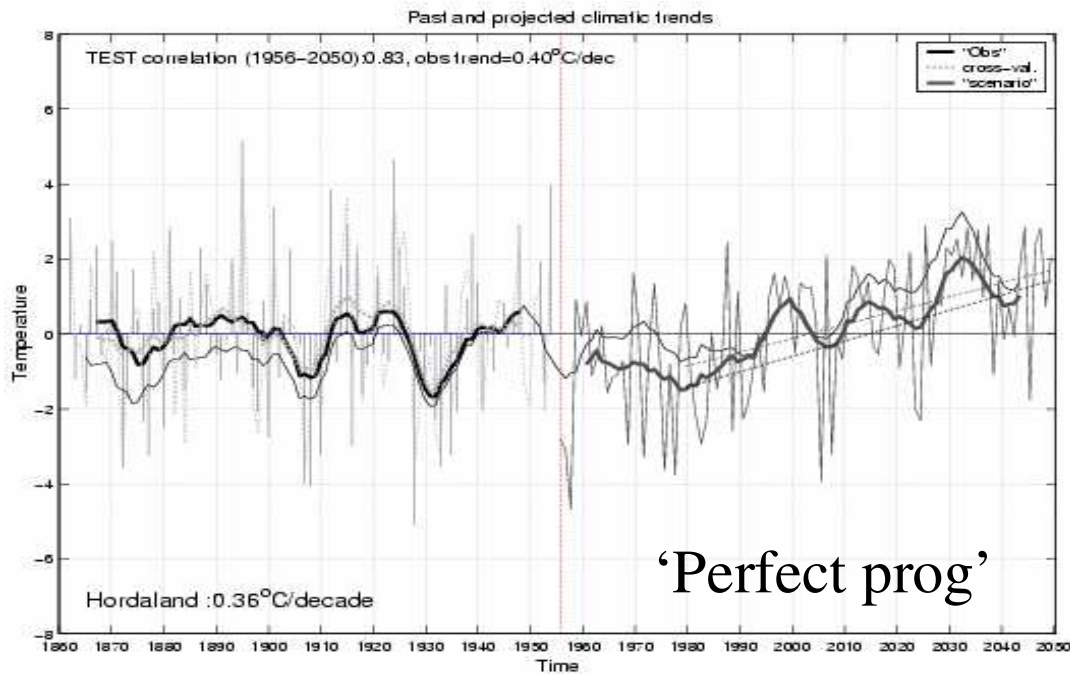
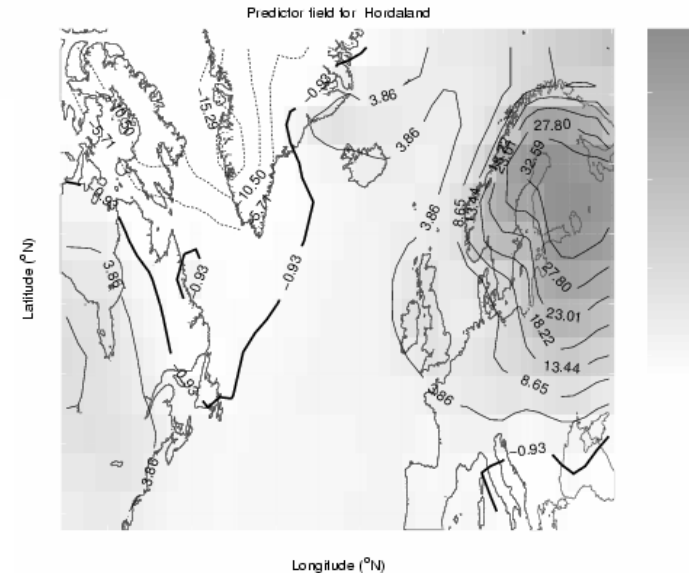
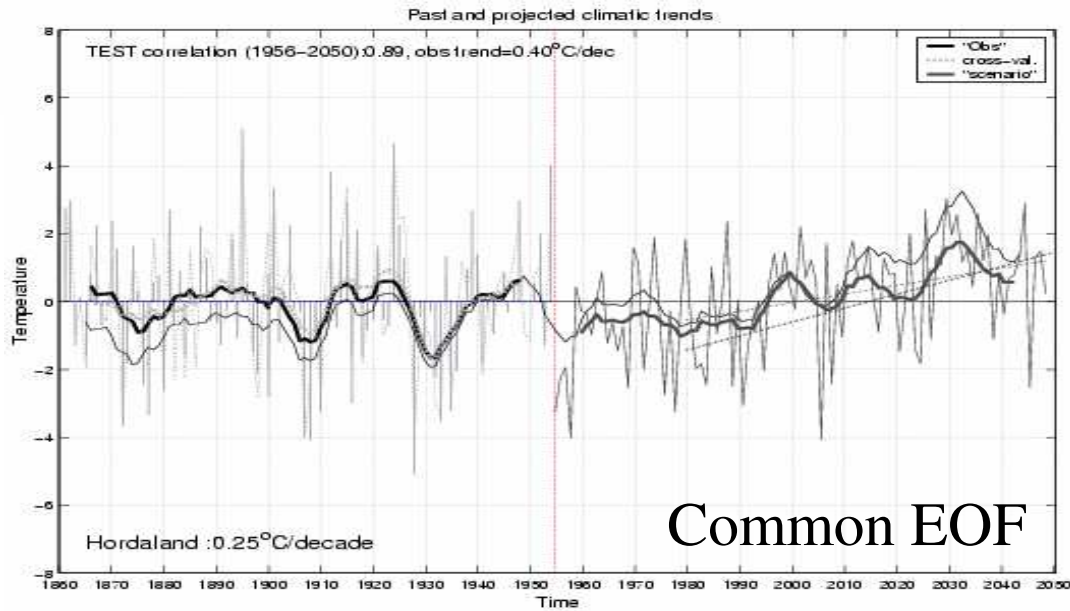
U: spatial pattern common
Σ: 'Eigenvalues' (variance)
V: time series describing the loadings (principal components)

Mathematically identical to Empirical orthogonal functions (EOF).

Methods



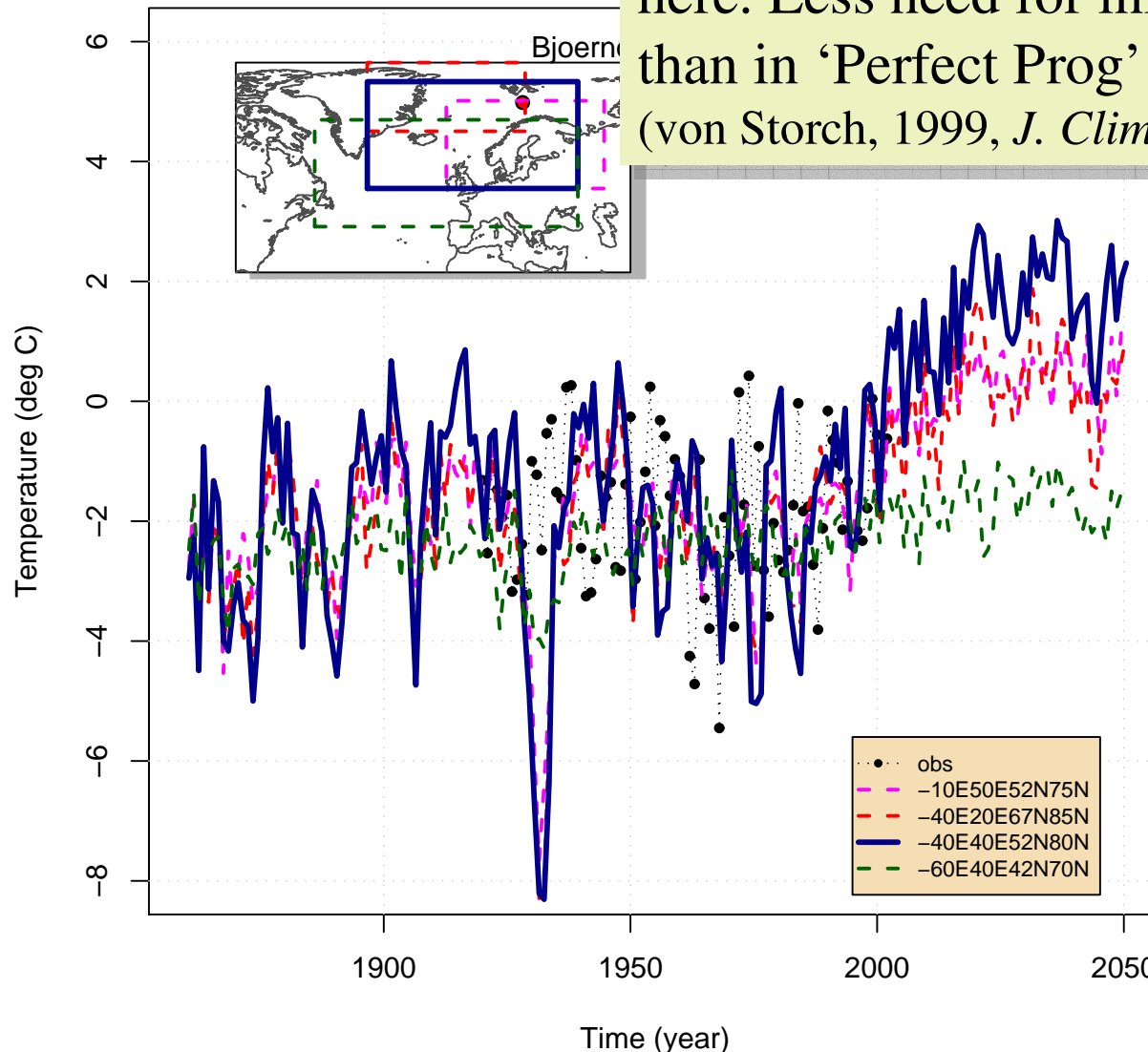
Example of downscaling: 'Perfect prog' & common EOF



Choice of domain can affect your results

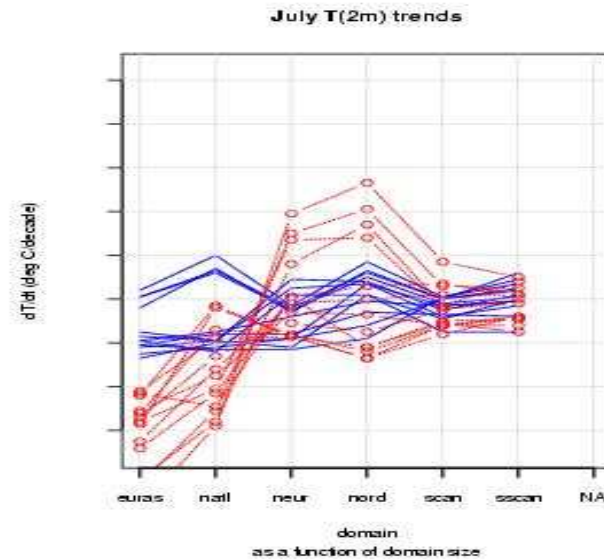
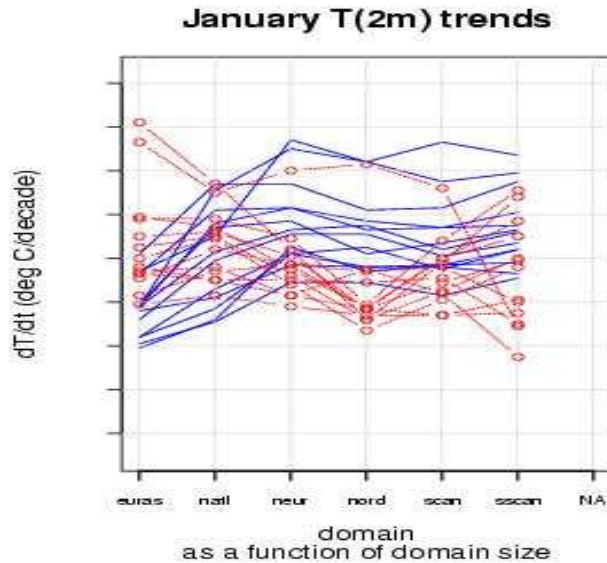


Annual mean temperature

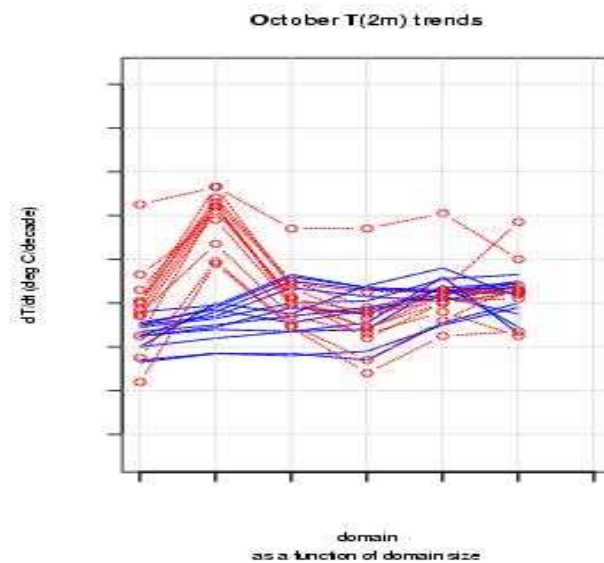
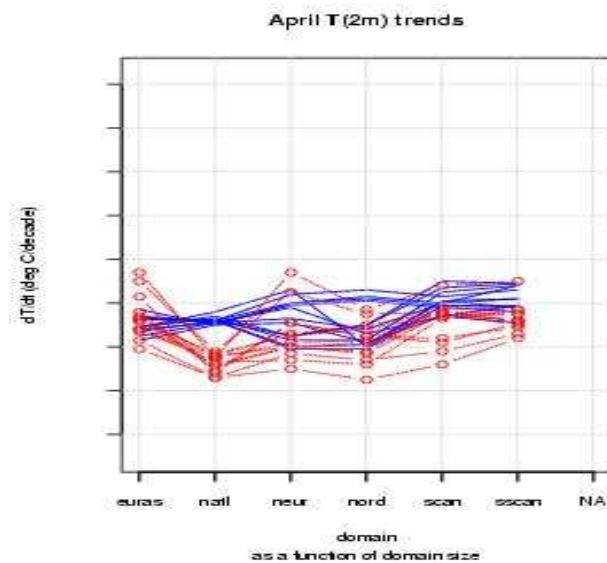


No 'inflation' has been used here. Less need for inflation than in 'Perfect Prog' approach (von Storch, 1999, *J. Clim*, 12, 3505)

Experiment: downscaling using a set of different predictor domains. Check robustness (flat structure).



Common EOF
'Perfect prog'





Empirical Orthogonal Functions (EOFs) and Principal Component Analysis (PCA).

Eigenvectors of the co-variance matrix:

‘S-mode’ and ‘T-mode’

$$[\vec{x}_1, \vec{x}_2, \dots, \vec{x}_m] \rightarrow \mathbf{X}$$

Variance-covariance matrix \mathbf{S} of \mathbf{X}' is $1/(n-1) \mathbf{X}'^T \mathbf{X}$ where

$$\mathbf{X}' = \mathbf{X} - \bar{\mathbf{X}}$$

$$\mathbf{S} \vec{e} = \lambda \vec{e} \text{ (Eigenfunctions)}$$

$$\vec{x}' = \mathbf{E} \vec{u}$$

$$\vec{u}_m = \vec{e}_m^T \vec{x}'$$

$$\vec{e}_i^T \vec{e}_j = \delta_{ij}$$

$$[\vec{e}_1, \vec{e}_2, \dots, \vec{e}_n] \rightarrow \mathbf{E}$$

Singular Vector Decomposition (SVD)

$$\mathbf{X} = \mathbf{U} \mathbf{\Sigma} \mathbf{V}^T$$

Literature:

Wilks, D.S. (1995) Statistical Methods in the Atmospheric Sciences. *Academic Press*

Press W.H., Flannery B.P., Teukolsky S.A., & Vetterling W.T..(1989) Numerical Recipes, *Cambridge*

Preisendorfer R.W. (1989) Principal Component Analysis in Meteorology and Oceanology, *Elsevier Science Press*

Regression & other Statistical models on Relationships



Regression:

single & mul

Least

multivariate

Proje

$$\vec{y} = a$$

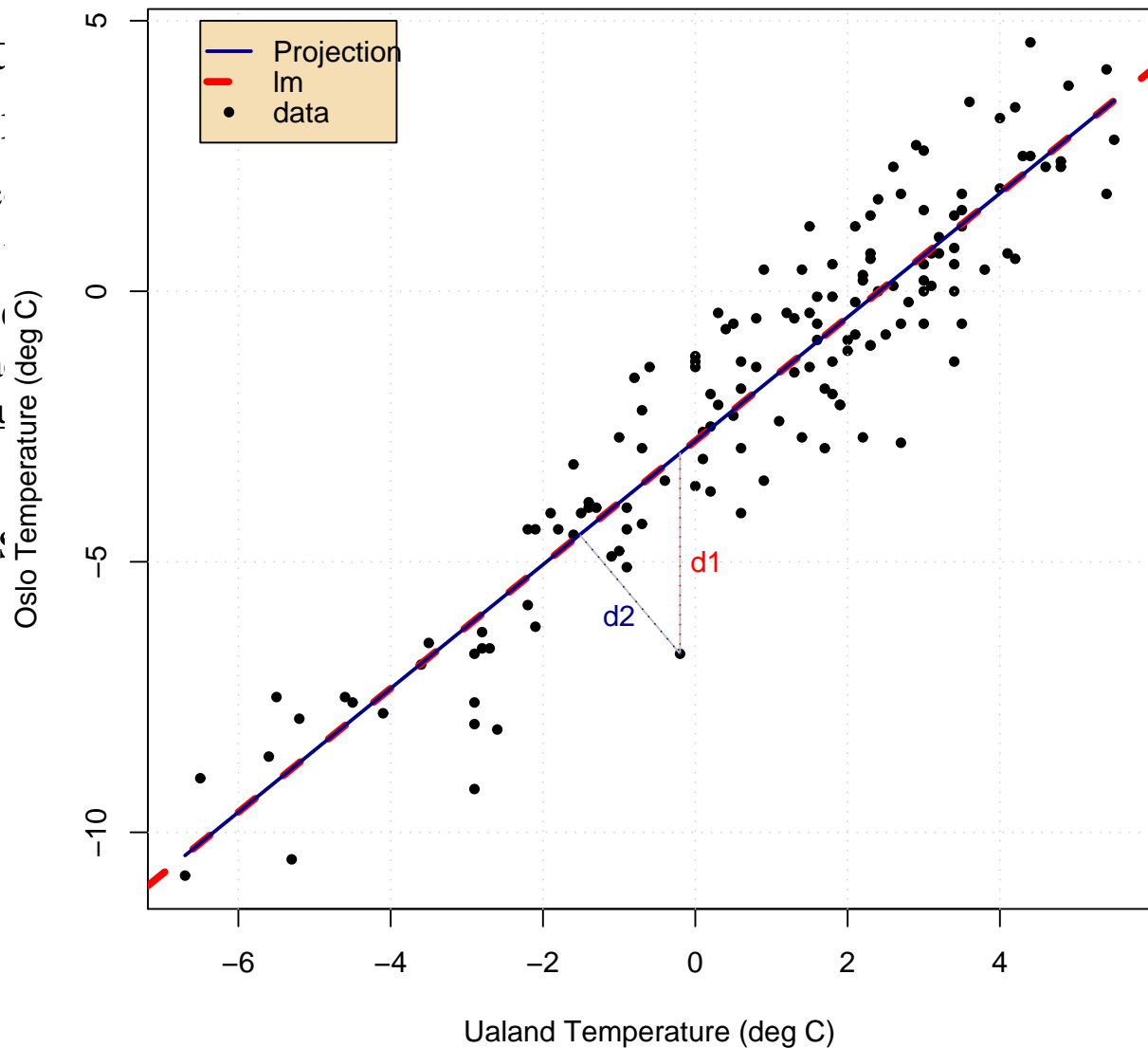
Strang,

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Mean Land Surface Temperatu

linear models

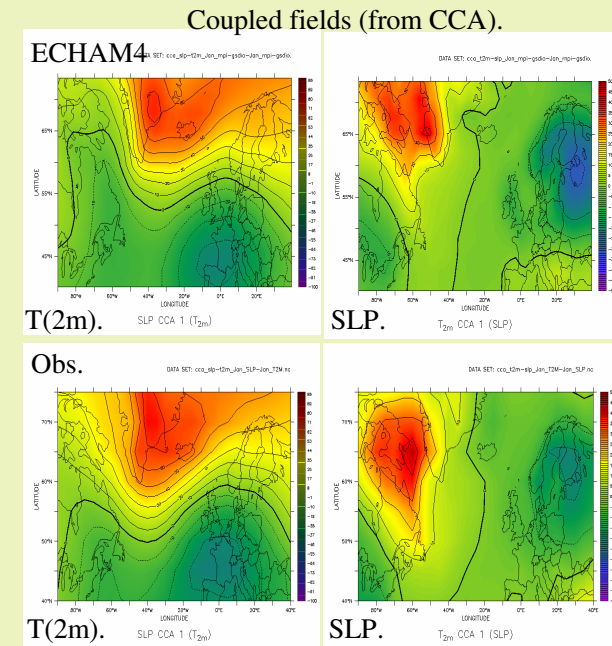
Example: Im versus projection



Canonical Correlation Analysis (CCA) classical & Barnett-Preisendorfer

Bretherton et al. (1992) An Intercomparison of
Coupled Patterns in Climate Data, *J. Clim.*, **5**, 541.

Benestad (1998) CCA applied to Statistical Downscaling
Monthly Mean Land Surface Temperatures: Model Document
28/98, pp.96



Find patterns with the
maximum correlation.

$$\mathbf{X}_1 = \mathbf{G} \mathbf{U}^T, \mathbf{X}_2 = \mathbf{H} \mathbf{V}^T$$

$$\mathbf{U}^T \mathbf{V} = \mathbf{L} \mathbf{M} \mathbf{R}^T = \mathbf{C}$$

Downscaling:

$$\widehat{\mathbf{X}}_1 = \mathbf{G} \mathbf{M} (\mathbf{H}^T \mathbf{H})^{-1} \mathbf{X}_2$$



Singular Vector Decomposition (SVD)

not to be confused with Singular Vector
Decomposition (SVD)

Benestad (1998) SVD applied to Statistical Downscaling for Prediction of Monthly Mean Land Surface Temperatures: Model Documentation, *DNMI Klima*, 30/98, pp. 38

$$\widehat{\mathbf{X}}_1 = \mathbf{G}_{\text{svd}} \mathbf{S}_{12} \mathbf{S}_{22}^{-1} \mathbf{H}_{\text{svd}}^T \mathbf{X}_2$$

Maximize co-variance (CCA maximizes correlation)

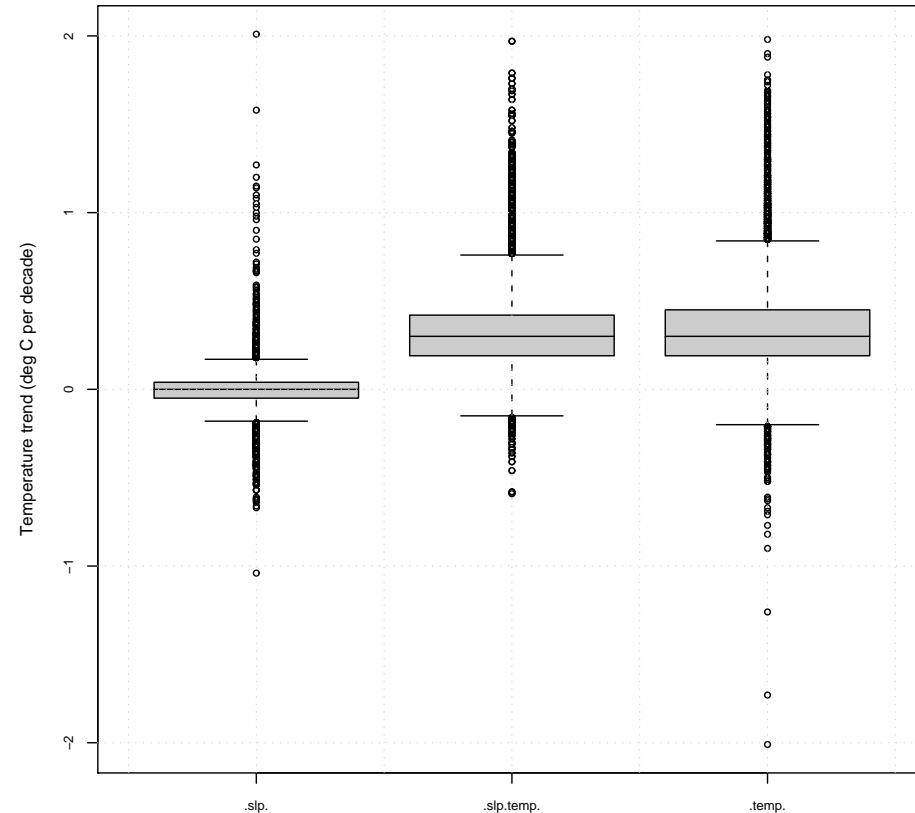
Other types of models...

Neural nets and Self-Organising Maps...



Which parameters as predictors?

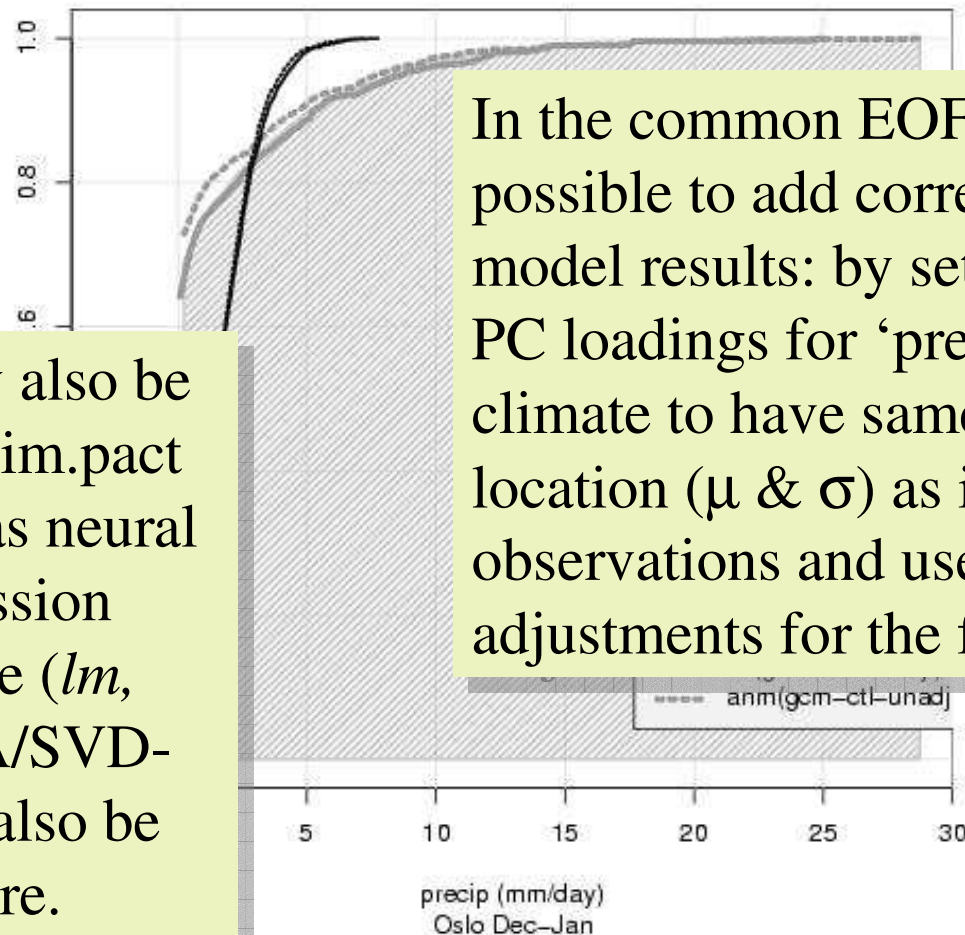
Strong & well-understood relationship (reflecting a physical mechanism)
Field that GCMs can skilfully reproduce
Parameters that carry the essential signal (e.g. a gradual global warming is not well-represented in SLP)





Choice of method: e.g. linear v.s. analog

Comparison between adjusted and unadjusted



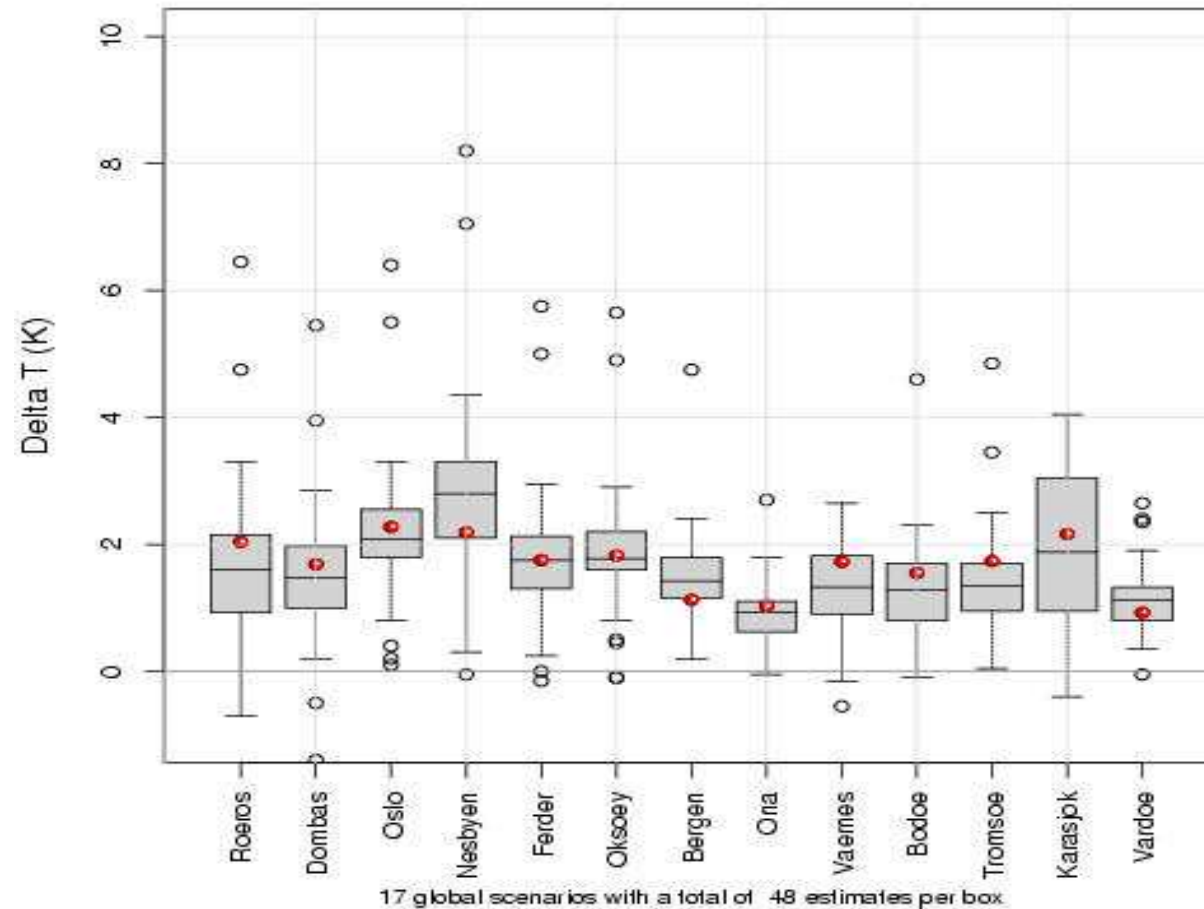
Other methods may also be incorporated into clim.pact in the future, such as neural nets. Various regression models are available (*lm*, *glm*, etc.), and CCA/SVD-based models may also be included in the future.

In the common EOF framework it is possible to add corrections to the model results: by setting PC loadings for ‘present-day’ climate to have same spread and location (μ & σ) as in the observations and use the same adjustments for the future.

Probing uncertainties through multi-model ensembles:
Spread caused by different model shortcomings, natural variability & different initialisation processes.



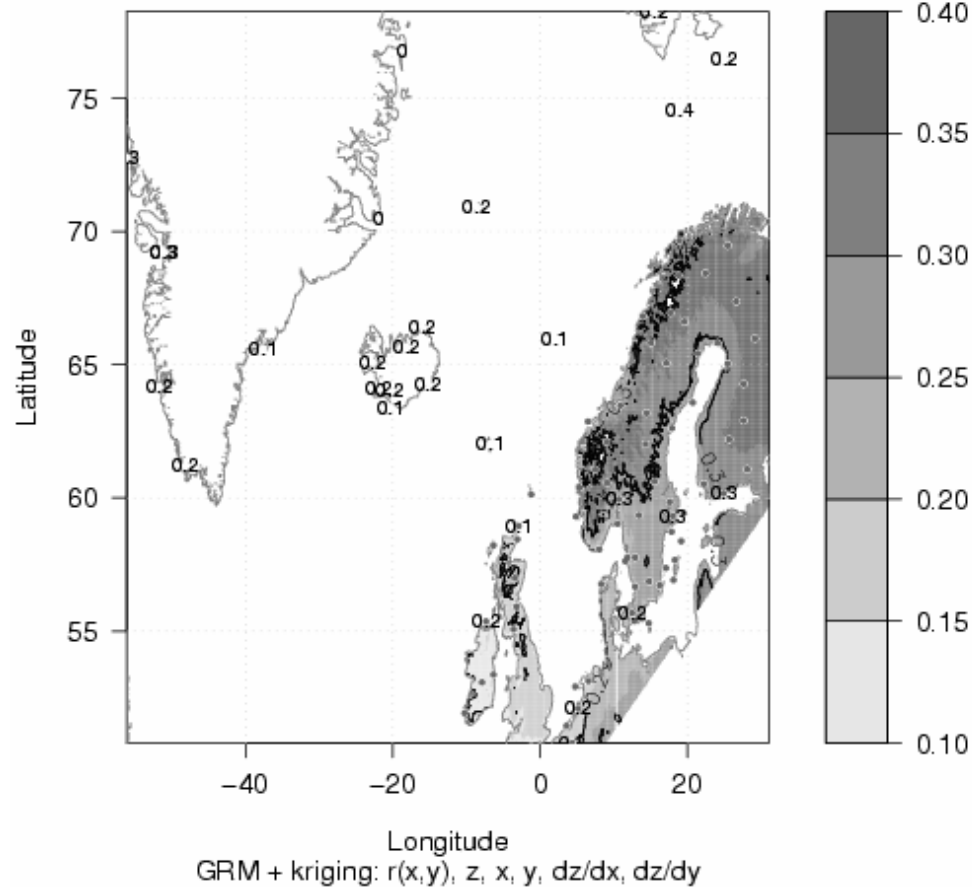
Jan scenario: temperature change (2000–2050)





Downscaling, ensembles & geographical distribution

Geographical model of temperature trend



OSLO annual temperature

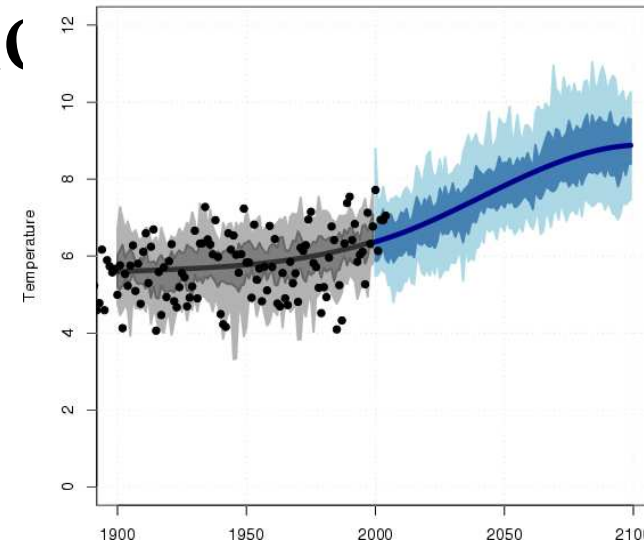


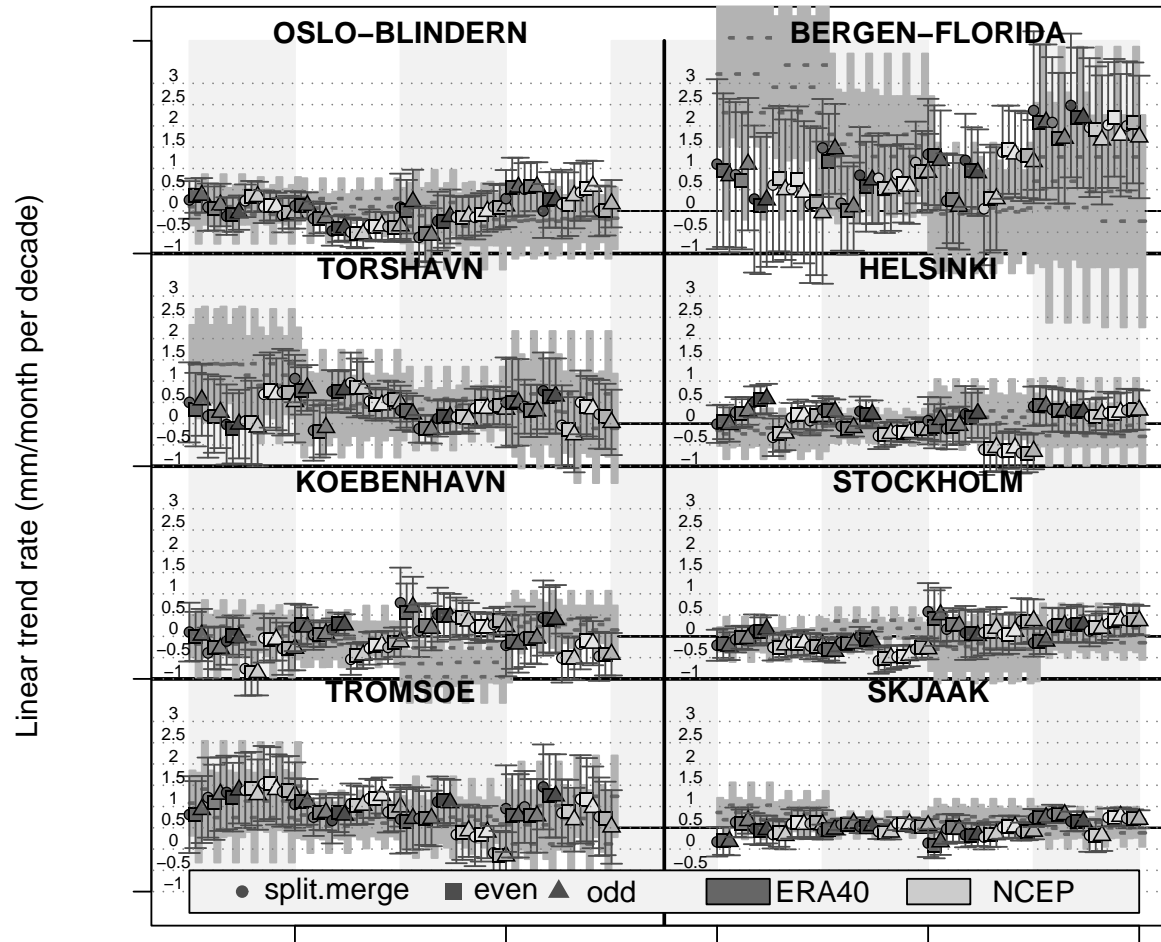
Table 3. Results from quality-control for GCMs with SRES A1b results for T(2m). The values are given in %.

GCM	f_s	f_m	n_{mco}	n_{pccgs}	n_{pcc1}	D	f_{var}	r_{sv}	n_m	N^*
CNRM-CM3	10.34	0.57	4.02	8.05	2.30	4.885	46.285	43.98	1.72	174
GFDL-CM2.0	4.60	0.57	0.57	7.47	2.30	4.885	20.405	43.68	0.57	174
GFDL-CM2.1	5.78	0.00	16.76	7.51	2.31	4.915	29.190	43.93	1.73	173
GISS-AOM	15.52	0.29	29.60	7.18	1.72	4.885	49.715	44.83	0.57	348
GISS-EH	3.26	0.19	6.32	6.90	2.30	4.885	34.005	41.38	0.57	822
GISS-ER	5.75	0.00	2.87	10.34	2.87	4.885	46.550	43.10	1.72	174
INM-CM3.0	17.82	0.57	0.57	7.47	2.30	4.885	41.380	44.83	7.47	174
IPSL-CM4	16.67	0.00	1.15	6.90	2.30	4.885	25.575	44.83	0.00	174
ECHAM5/MPI-OM	13.63	0.38	0.77	7.49	2.69	4.895	42.805	47.60	0.77	821
MRI-CGCM2.3.2	16.91	0.57	0.00	6.97	2.29	4.855	50.345	45.03	1.37	875
CCSMB	18.29	0.57	0.57	7.14	2.29	4.855	43.430	45.14	1.43	360
PCM	0.57	0.57	1.14	7.43	2.29	4.855	29.145	41.71	0.57	175
UKMO-HadCM3	23.43	0.57	1.14	6.86	2.29	4.855	41.715	45.71	0.00	175

Precipitation



Slope estimates

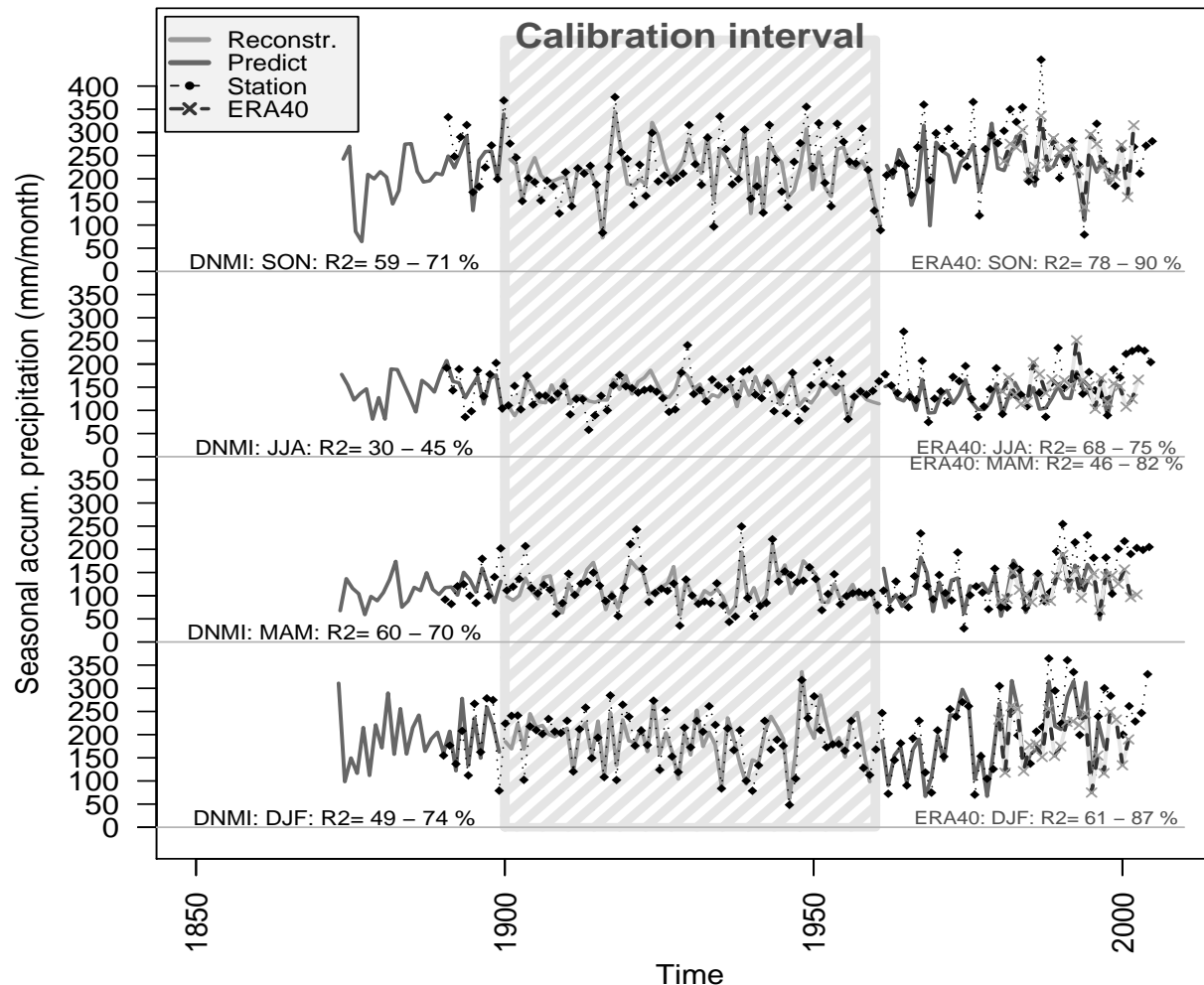


Predictors= prec , slp & mix



Validation of models

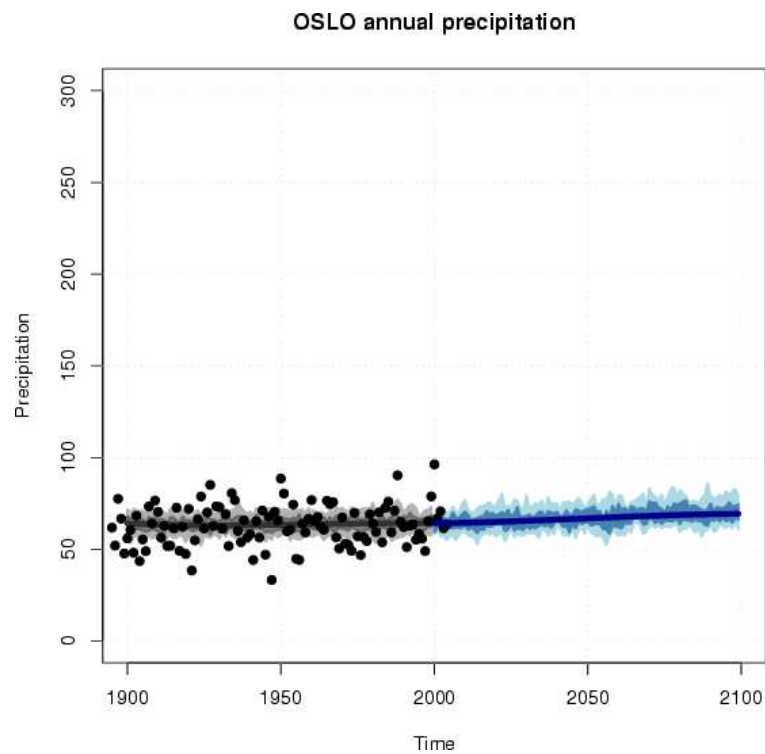
Bergen: reconstruction from gridded SLP



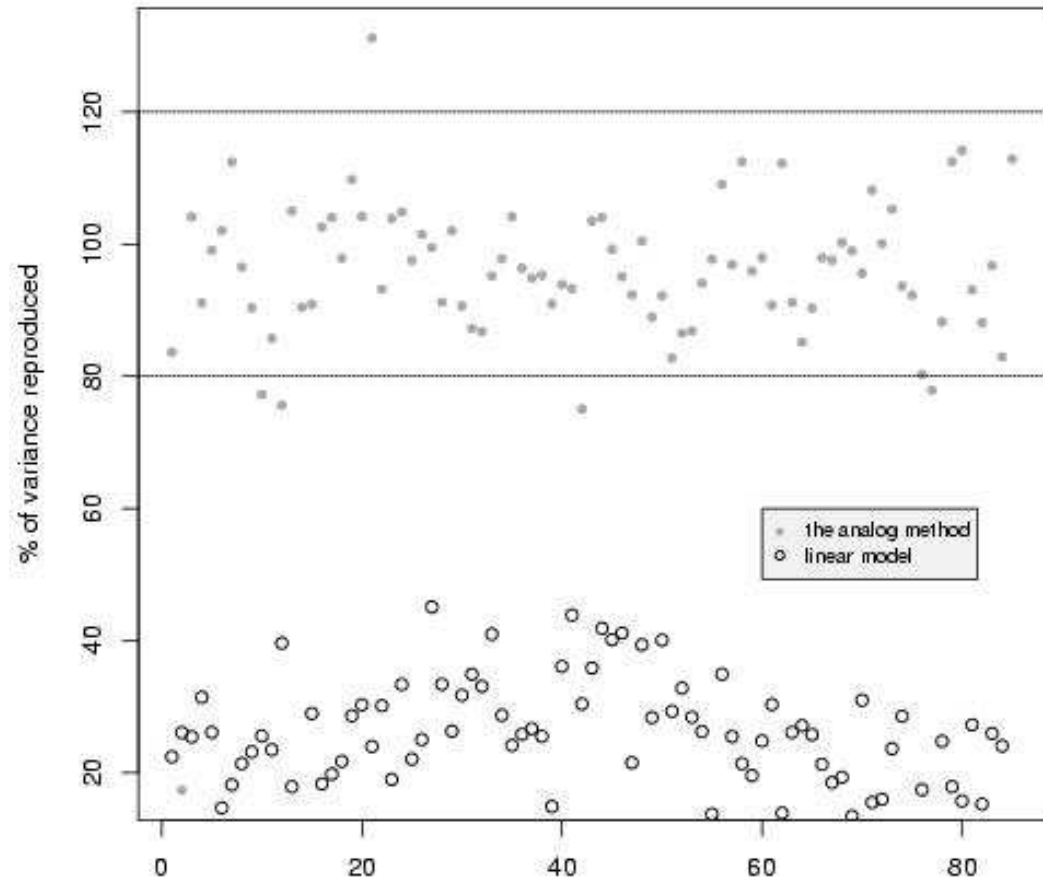
clim.pact: DNMI.slp (Benestad & Melsom, 2002, Clim. Res., Vol 23, 67–79)



When a fraction of the variance can be accounted



Change in variance between the scenario and the control intervals

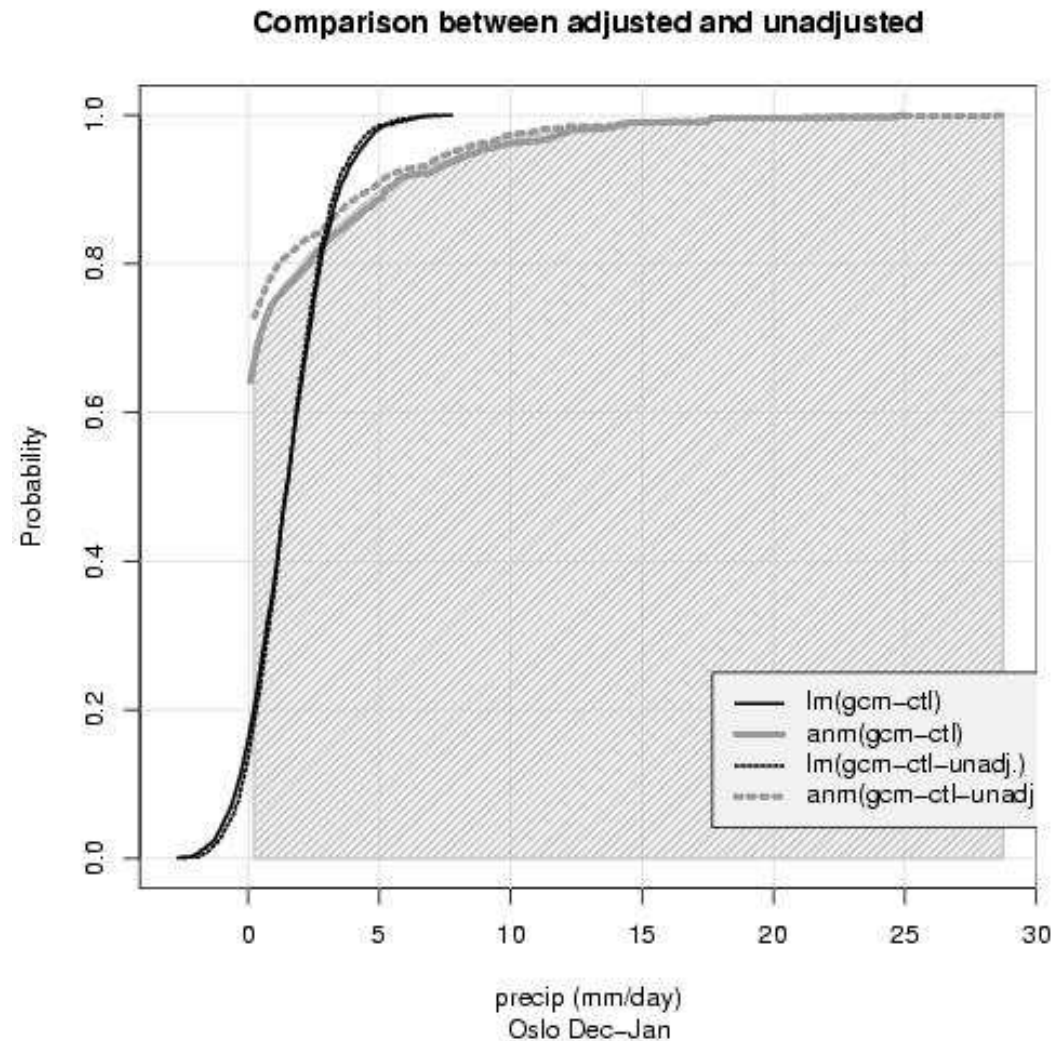


Predictor -> precipitation (mm) using era-15+psl+psl_SLP&PSL_5E12E-57N66N
Mean of variance for anm= 95.38%, for lm=25.65%



Daily rainfall

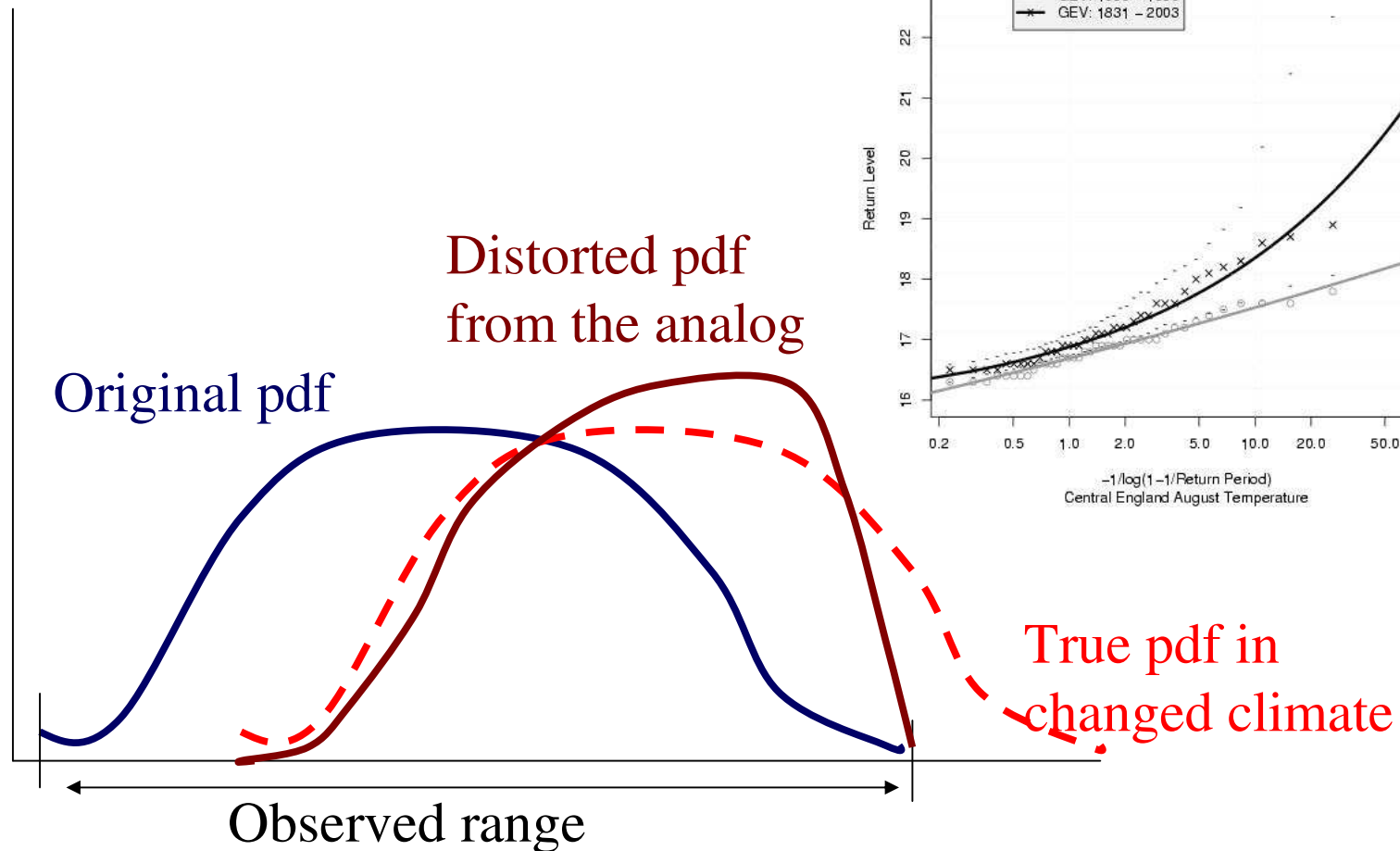
Linear (regression) models fail to give a good representation of the tails of the distribution. Other approach: **the analog model.**



The upper tail

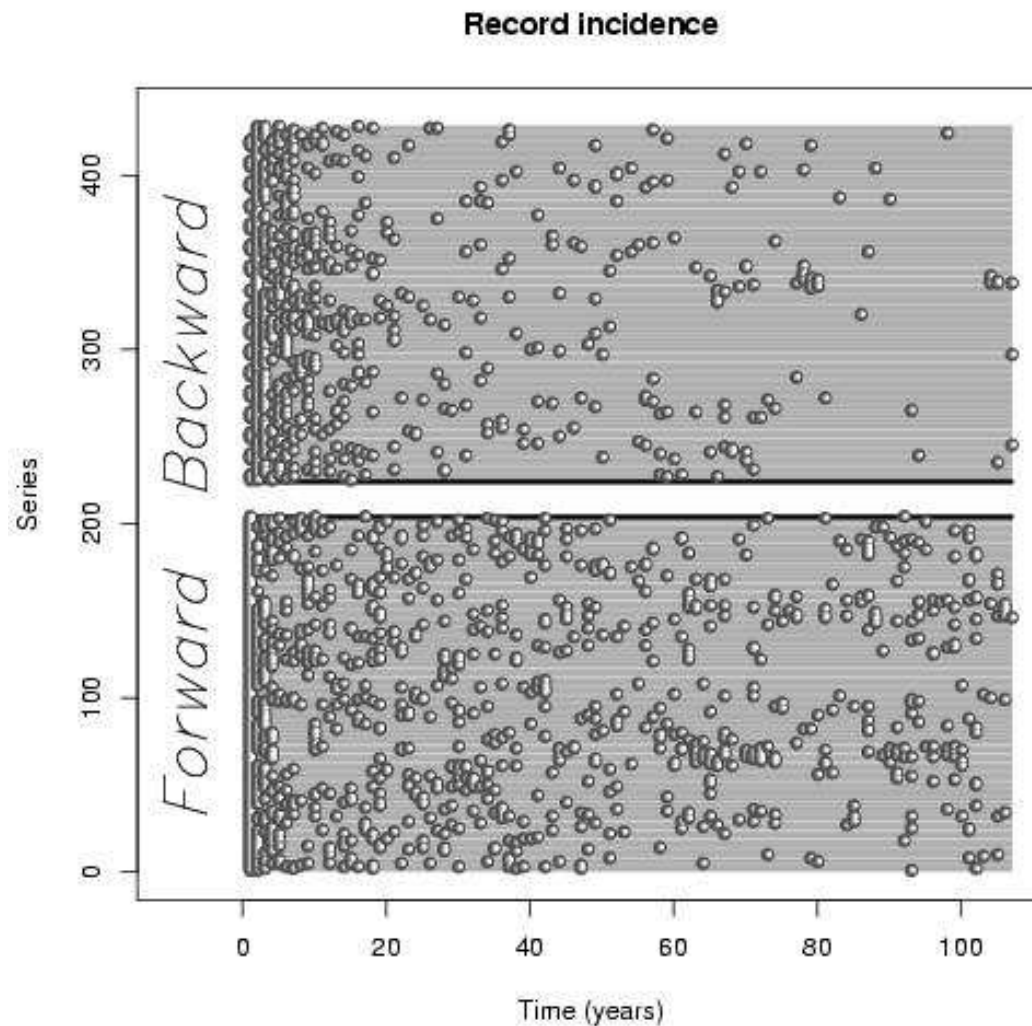


Analog model not representative for a situation where the upper tail of the distribution (pdf) is being

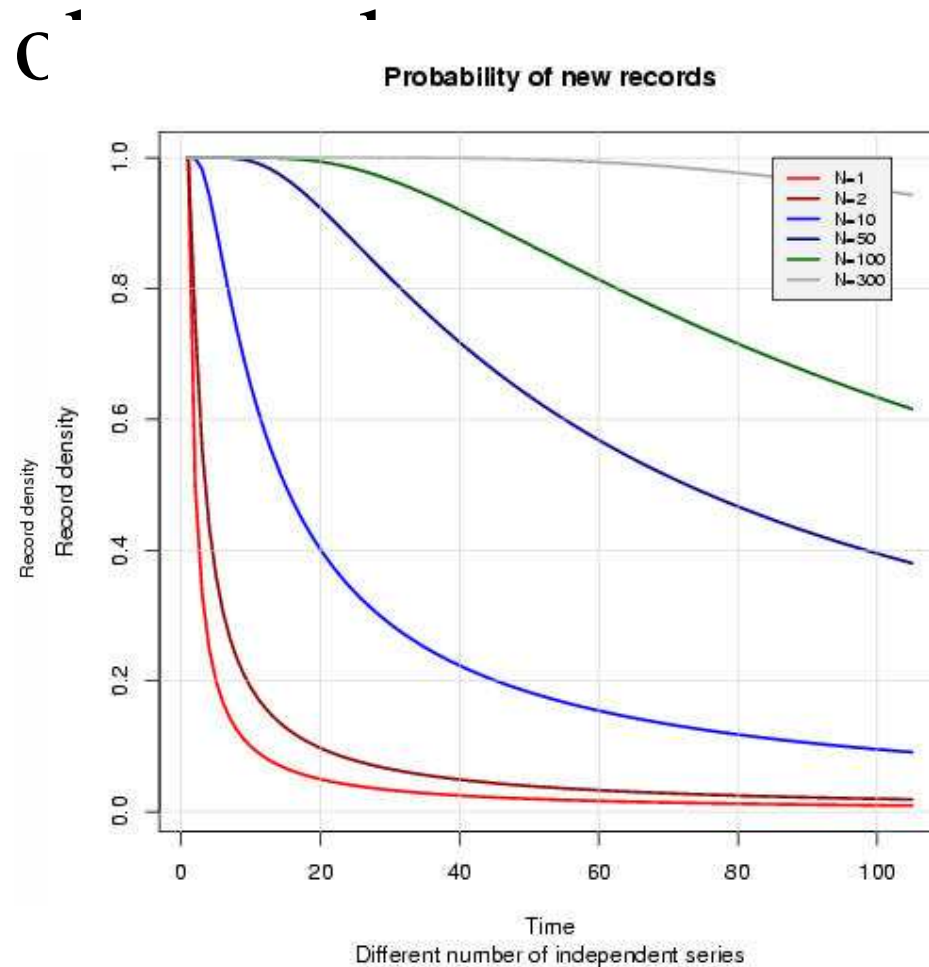




Record-events= values outside historical sample range



Caveat: the traditional analog model cannot predict values outside the



A random variable of rational numbers with independent and identical distribution (iid) has following property:

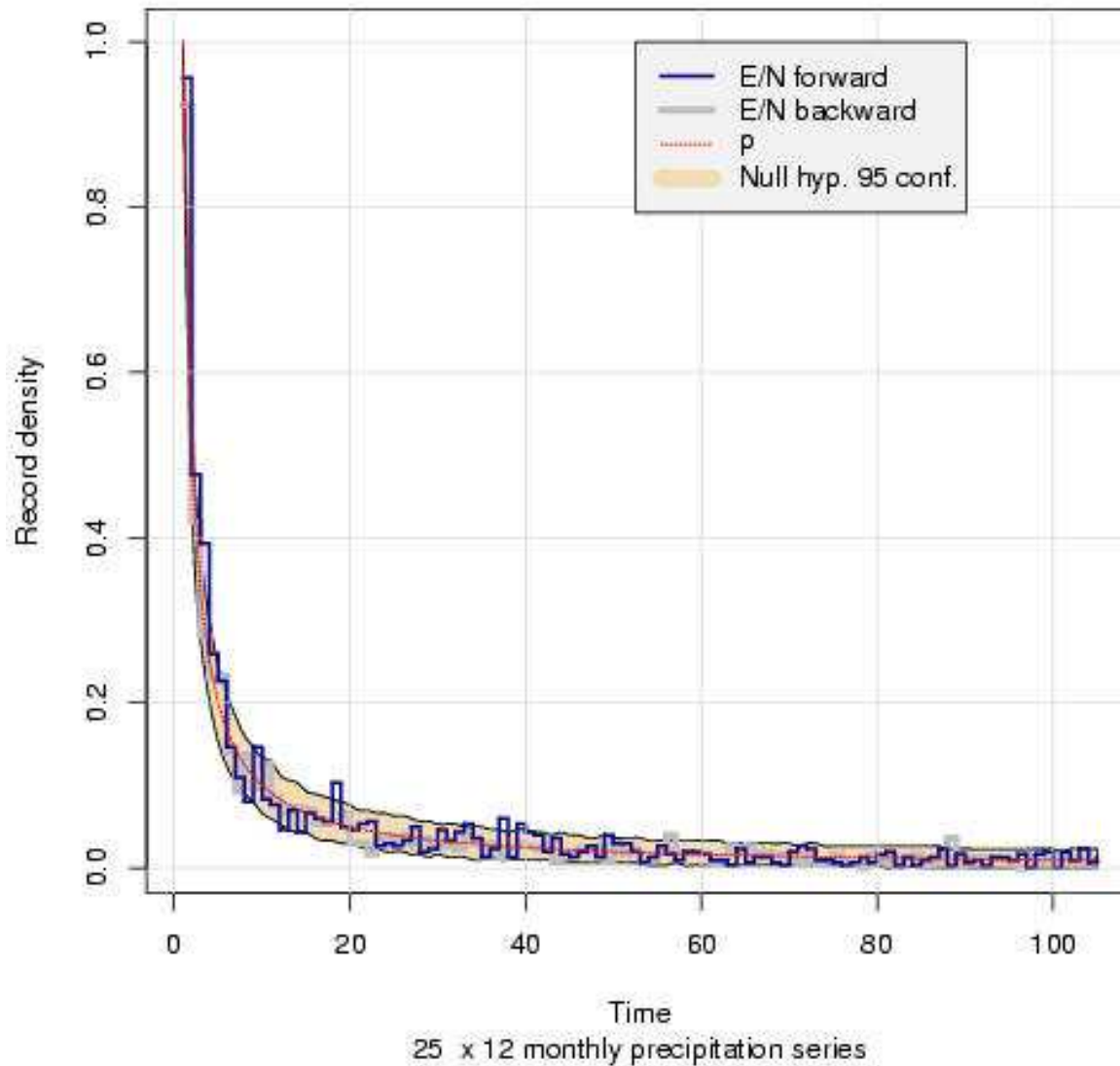
$$\Pr(n=\text{record}) = 1/n$$

(assuming no ties)



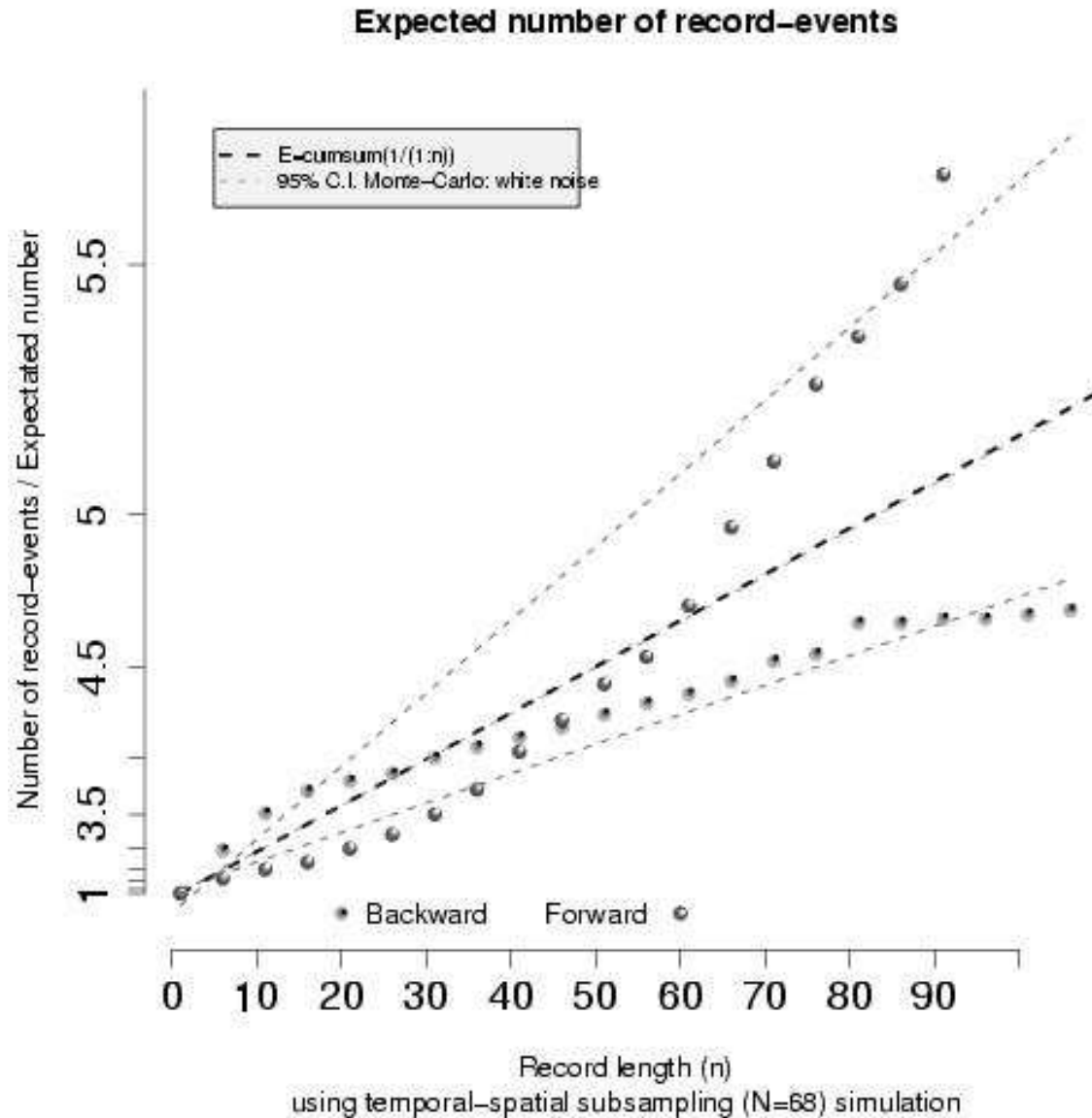
Record-event statistics

New records



Test of number of record-events:

iid



Re-cap of Uncertainties...



- Mean values - extremes are difficult
- Model shortcomings (systematic biases).
 - *Can be explored to a certain extent through multi-model ensembles.*
 - *Analysis, validations, & diagnostics (clim.pact)*
- Imperfect downscaling methods and choice of set-up.
 - *Can be explored through different choices and methods (eg clim.pact, SDSM, & nested models)*
 - *Validation exercises.*
- Natural variability, GCM initiation, & additional forcing factors.
 - *Ensembles of GCM experiments.*
- Emission scenarios.
 - *Multi-emission scenario ensembles.*



Possible solution

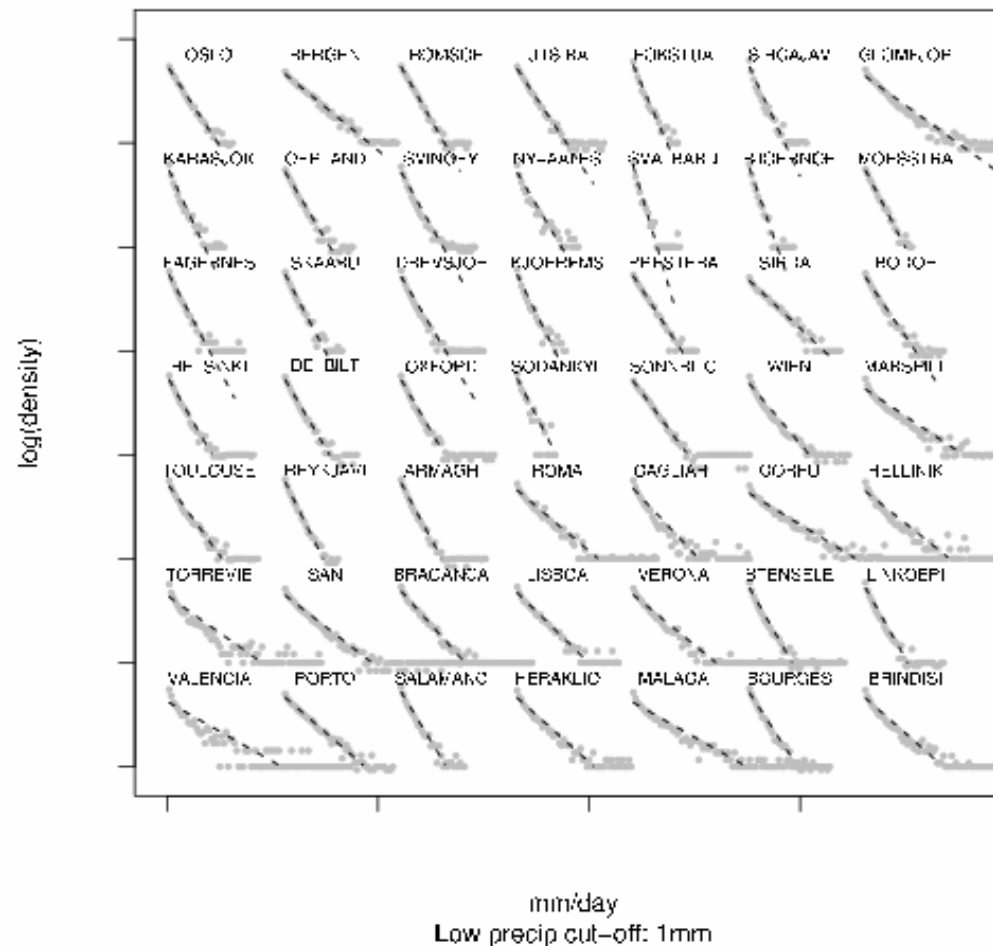
- Downscale the pdf?
- If the type of pdf shape is given (eg Gaussian, gamma, etc), and the parameters determining the exact structure of the pdf are systematically influenced by a set of conditions, then it may be possible to downscale the parameters for the pdfs.

Dependency of pdf on local conditions



For many situations, the linear-log relationships are approximately linear, indicating that an exponential law provides a reasonable fit. But, still some discrepancies in the upper tail.

Exp law: daily precipitation (1-order polynomial)





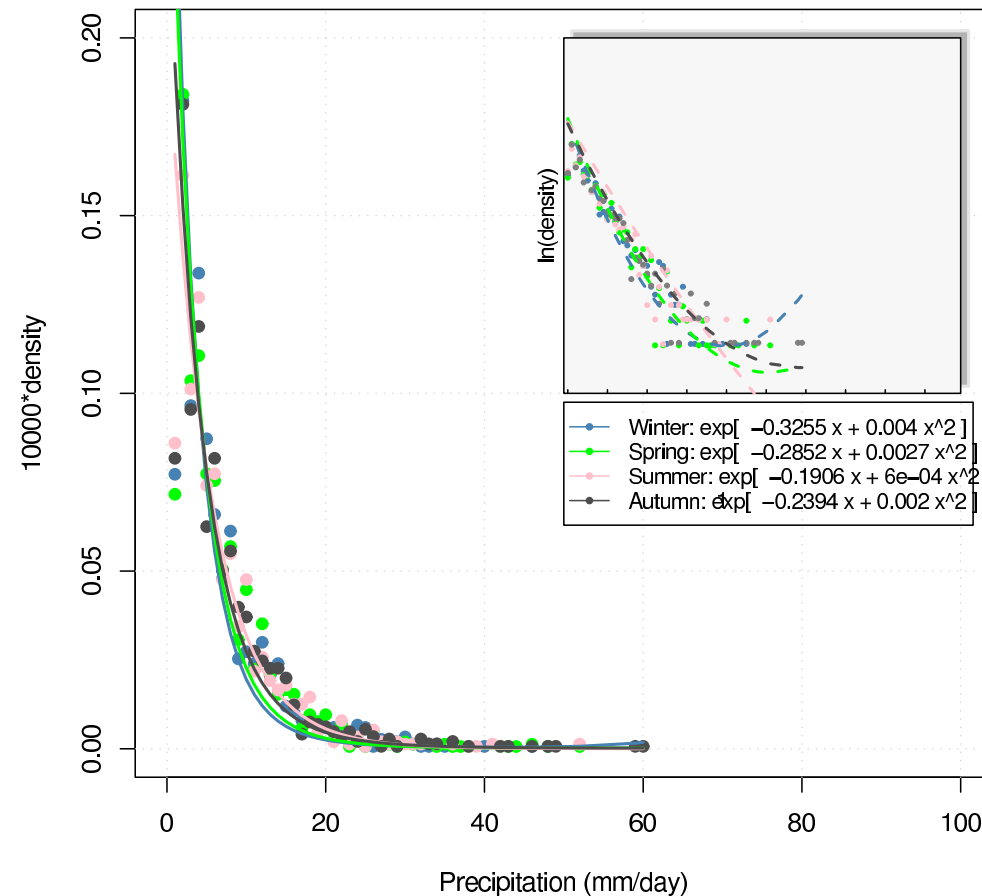
Exponential law: simpler than the gamma distribution

pdf:

$$p(x) = -m \exp\{-mx\}$$

m varies with local mean temperature and precipitation

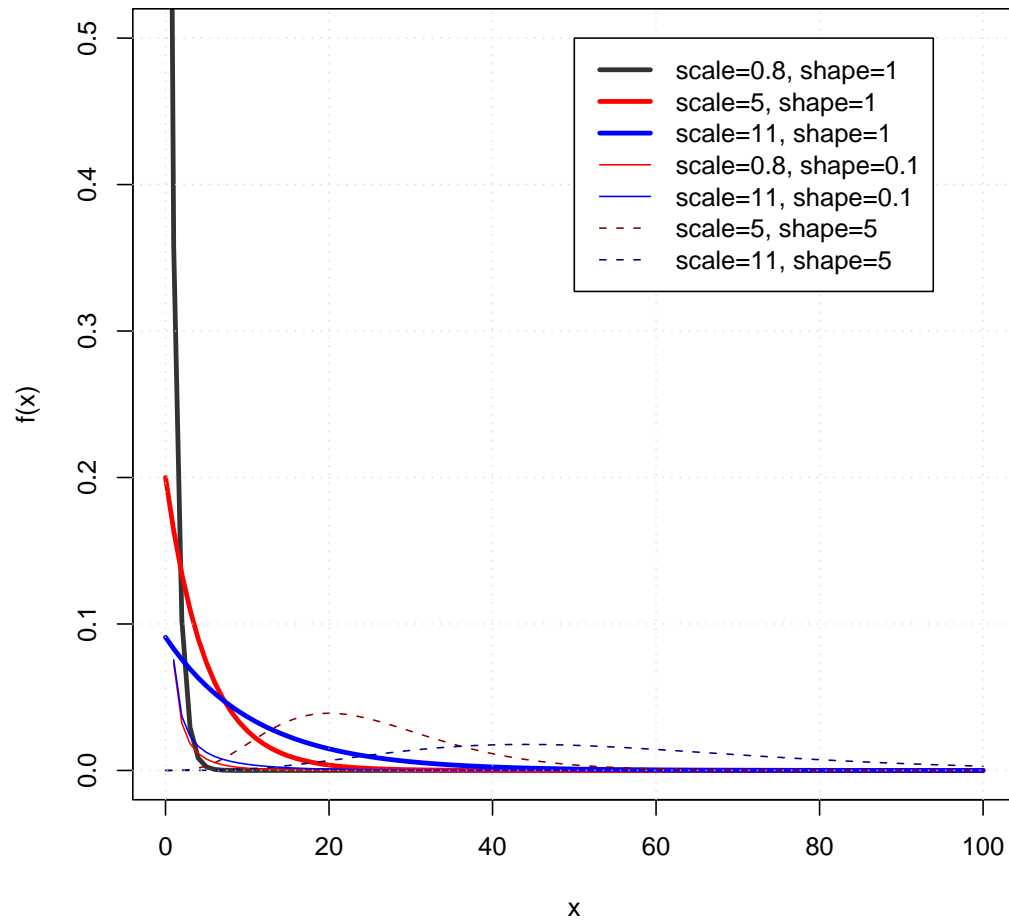
OSLO – BLINDERN



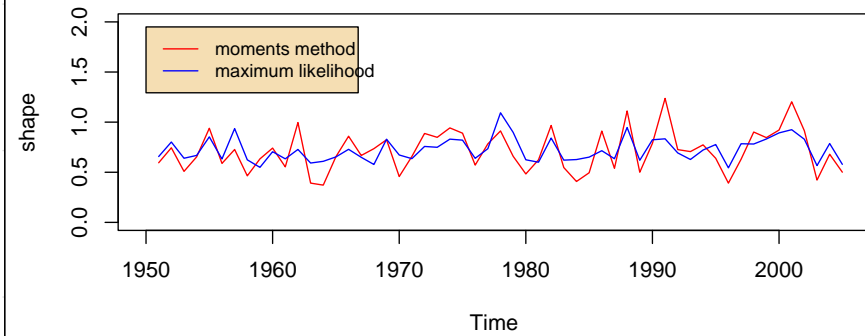


Ways of modelling the tails

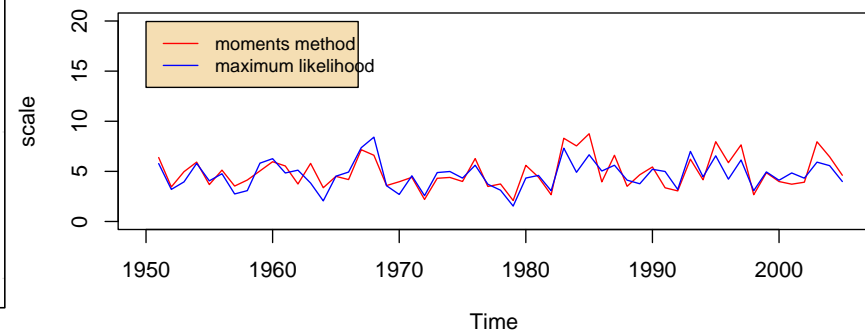
Gamma functions



OSLO – BLINDERN DJF

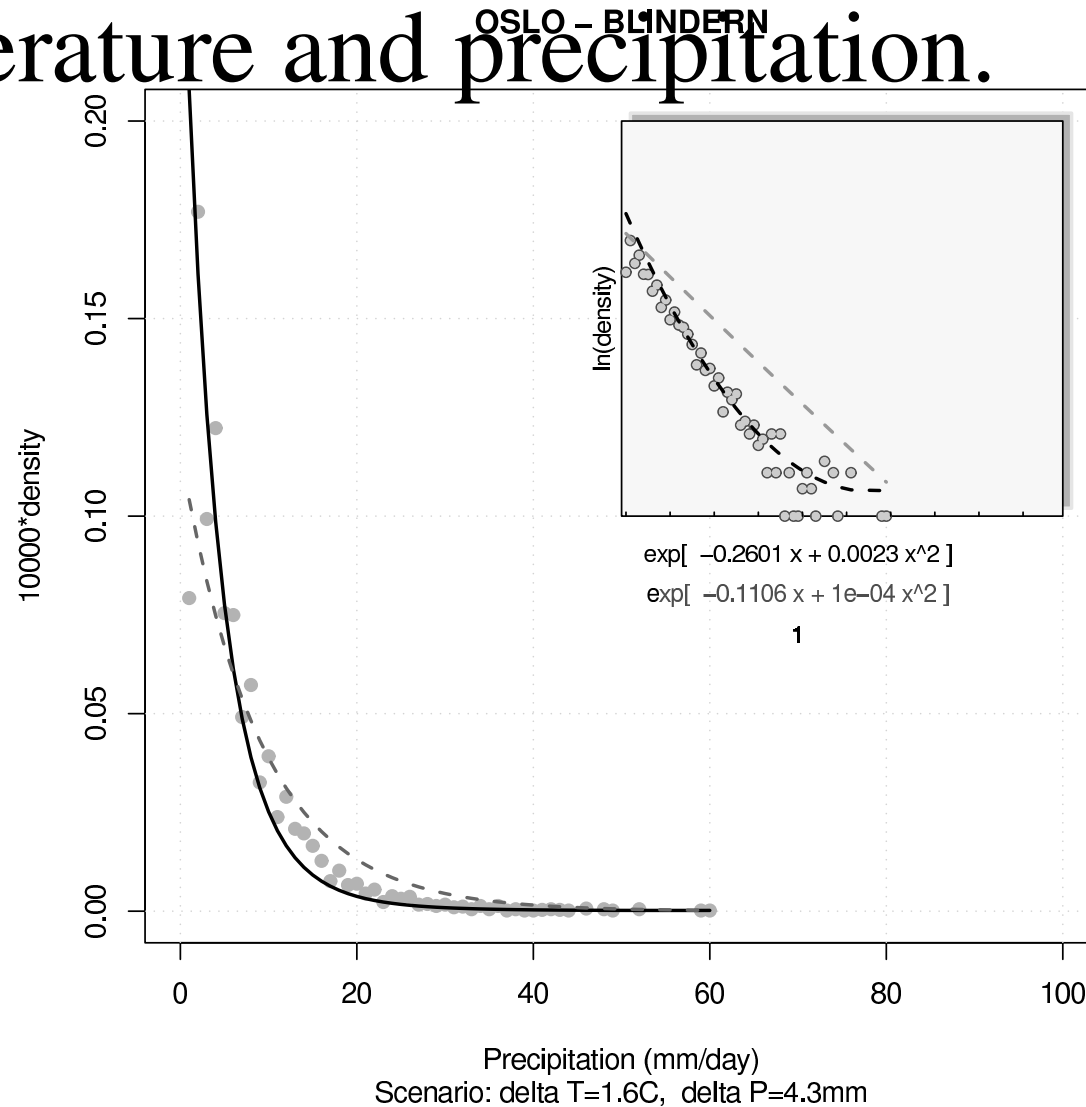


OSLO – BLINDERN DJF





Scenario for the future: derived from downscaled changes in the mean temperature and precipitation.



Exponential law: simple expression for unner nercentiles

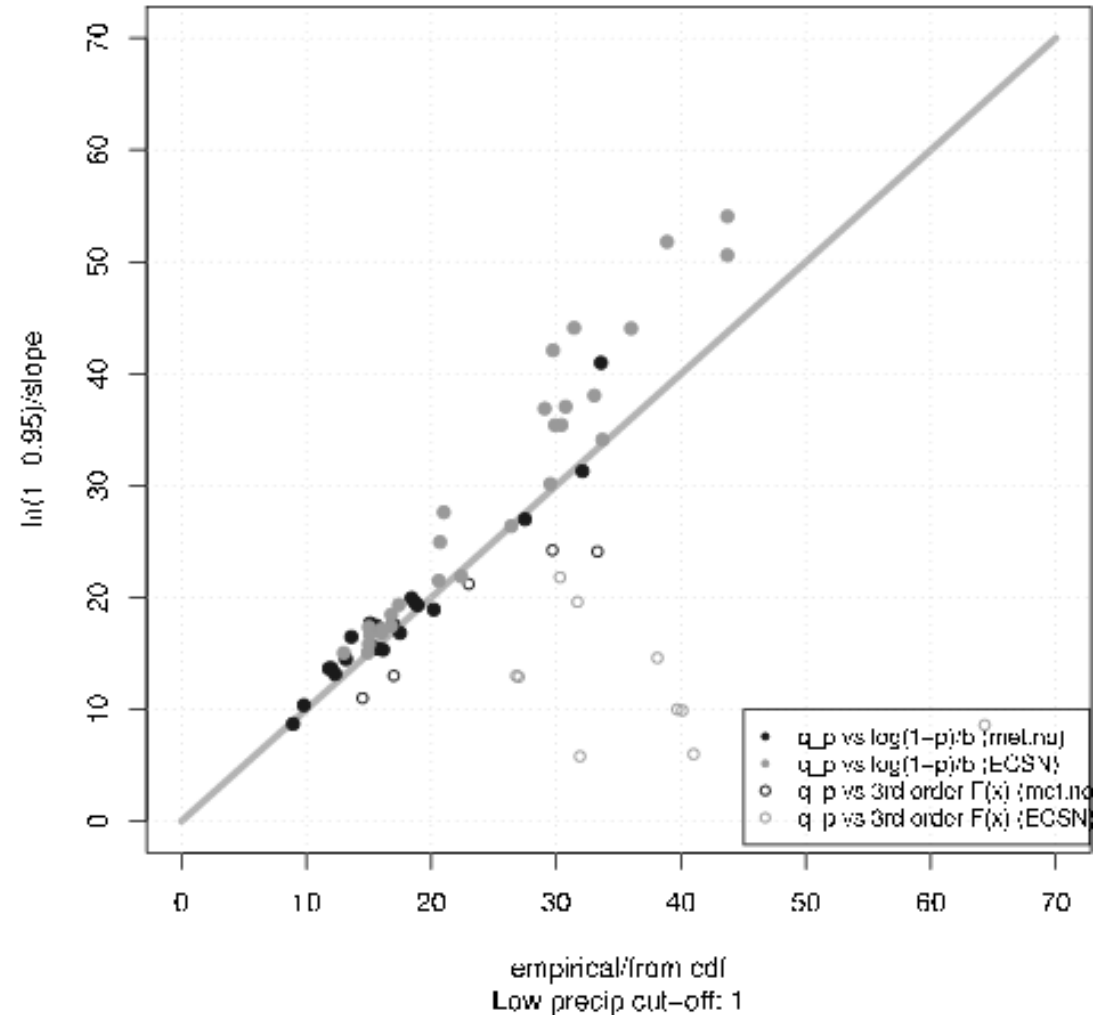


0.95 percentile: empirical v.s. theoretical

$$Q_p = \log(1-p)/m$$

$$M = -1/m$$

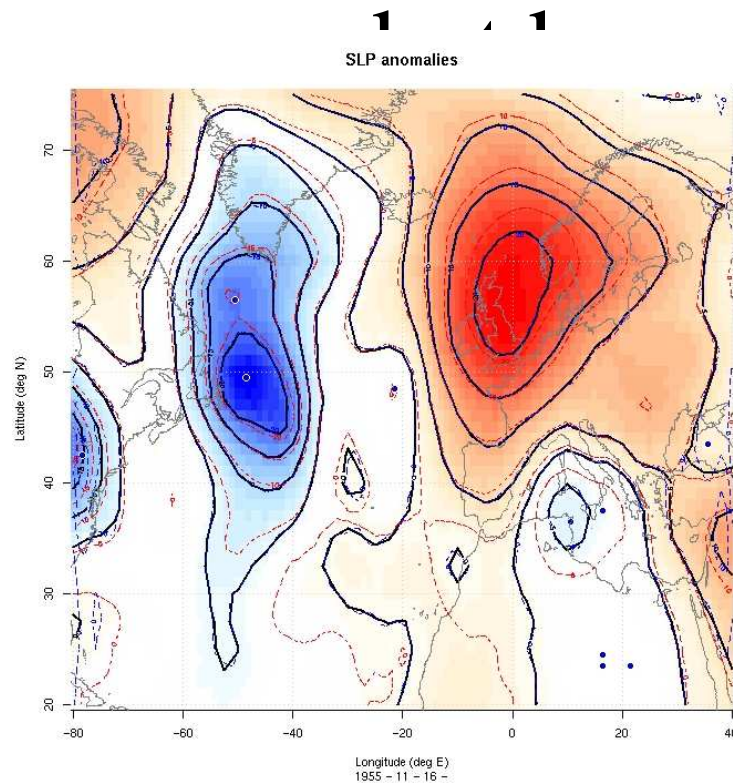
$$S = -2/m^2$$



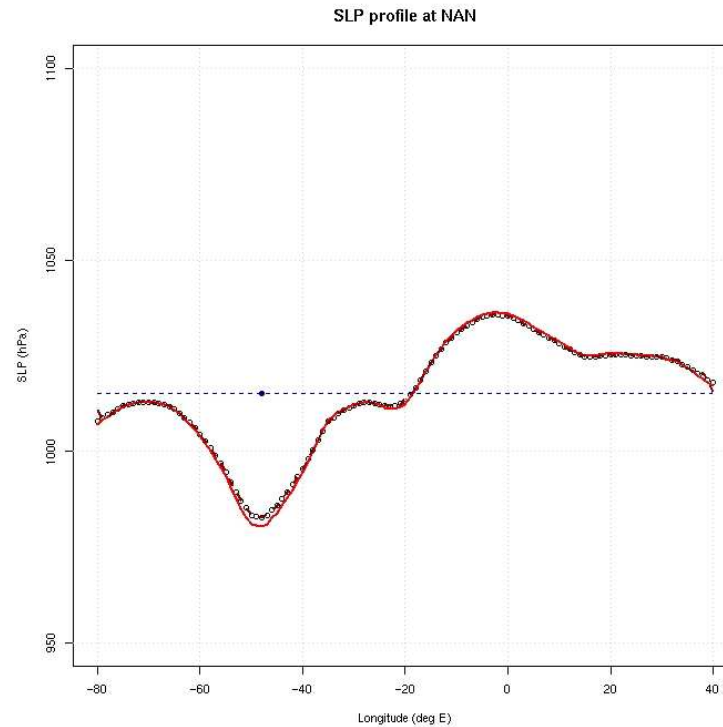


cyclones

(usually, we are not particularly interested in extremes because they



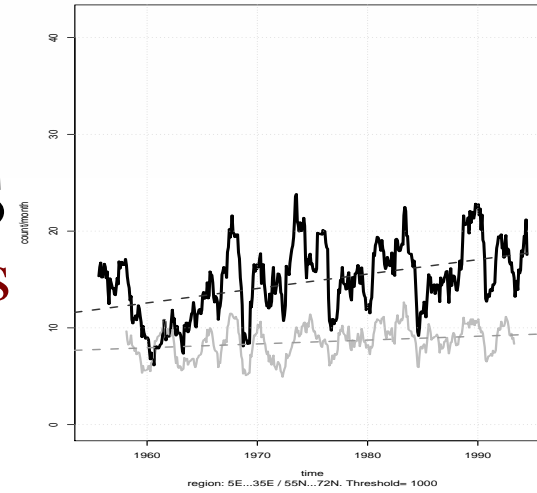
se



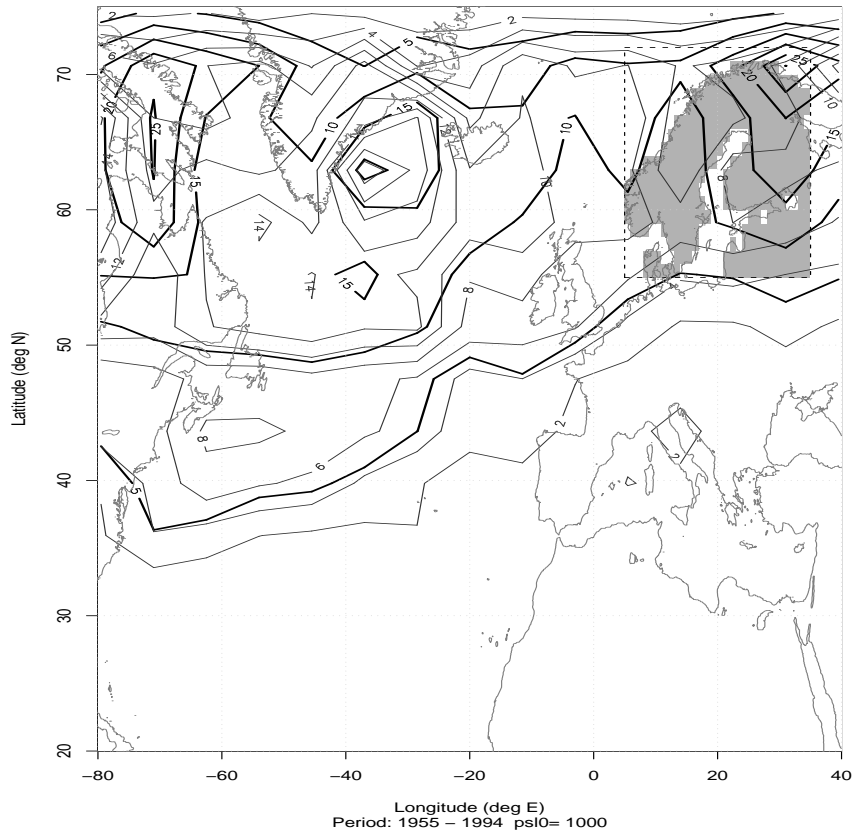
Downscaling cyclones Validation: analysis of observations



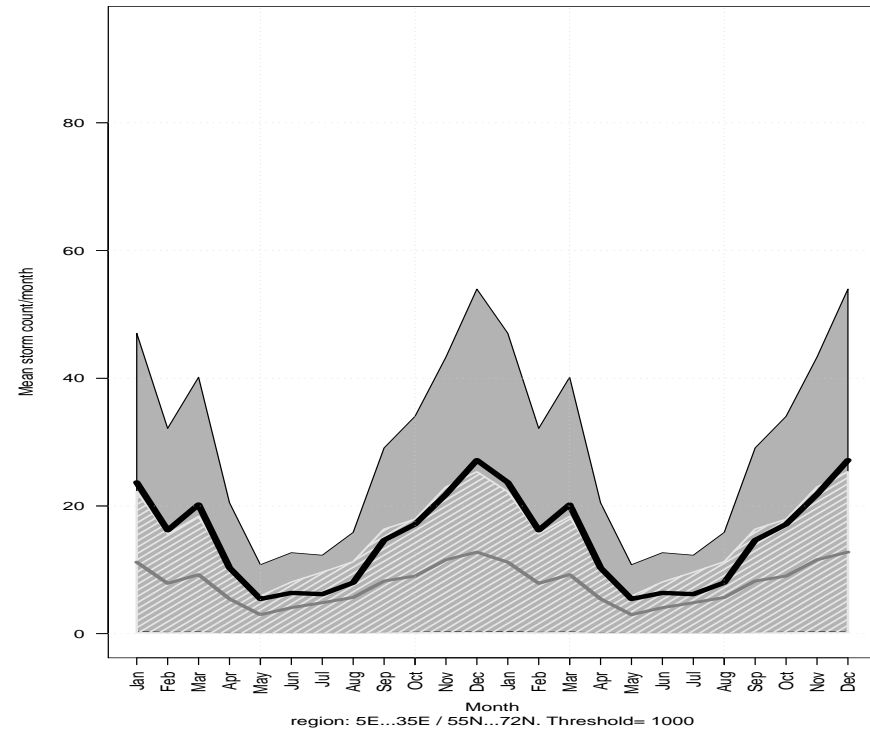
Cyclone count 1955 - 1994



Cyclone count per year



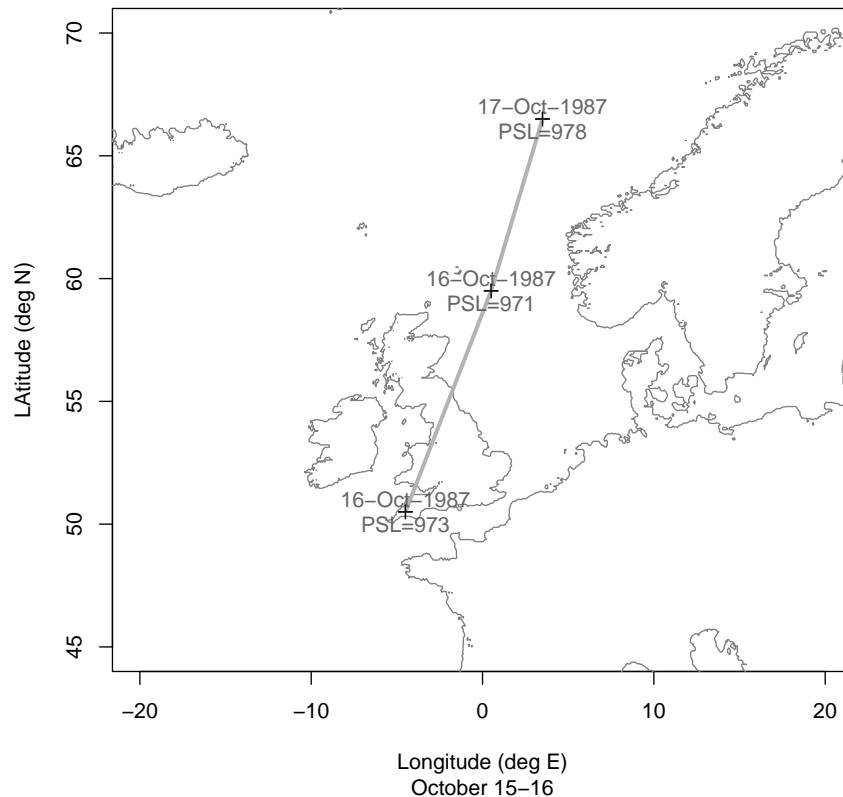
seasonal cyclone variability



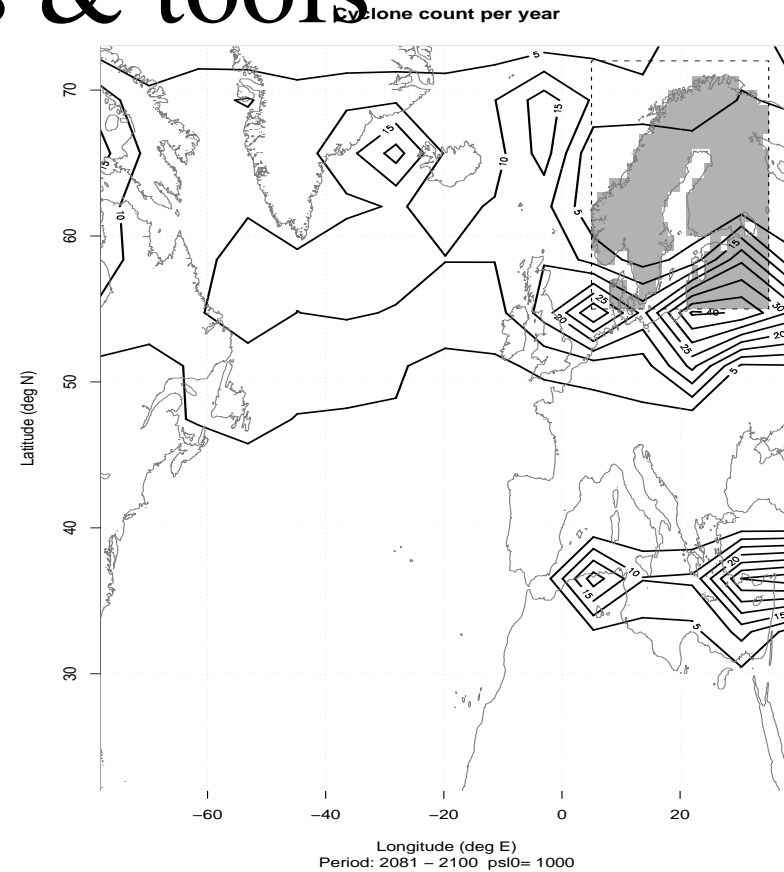


Testing the models & tools

The Great 1987 Storm



Reproduces known storms

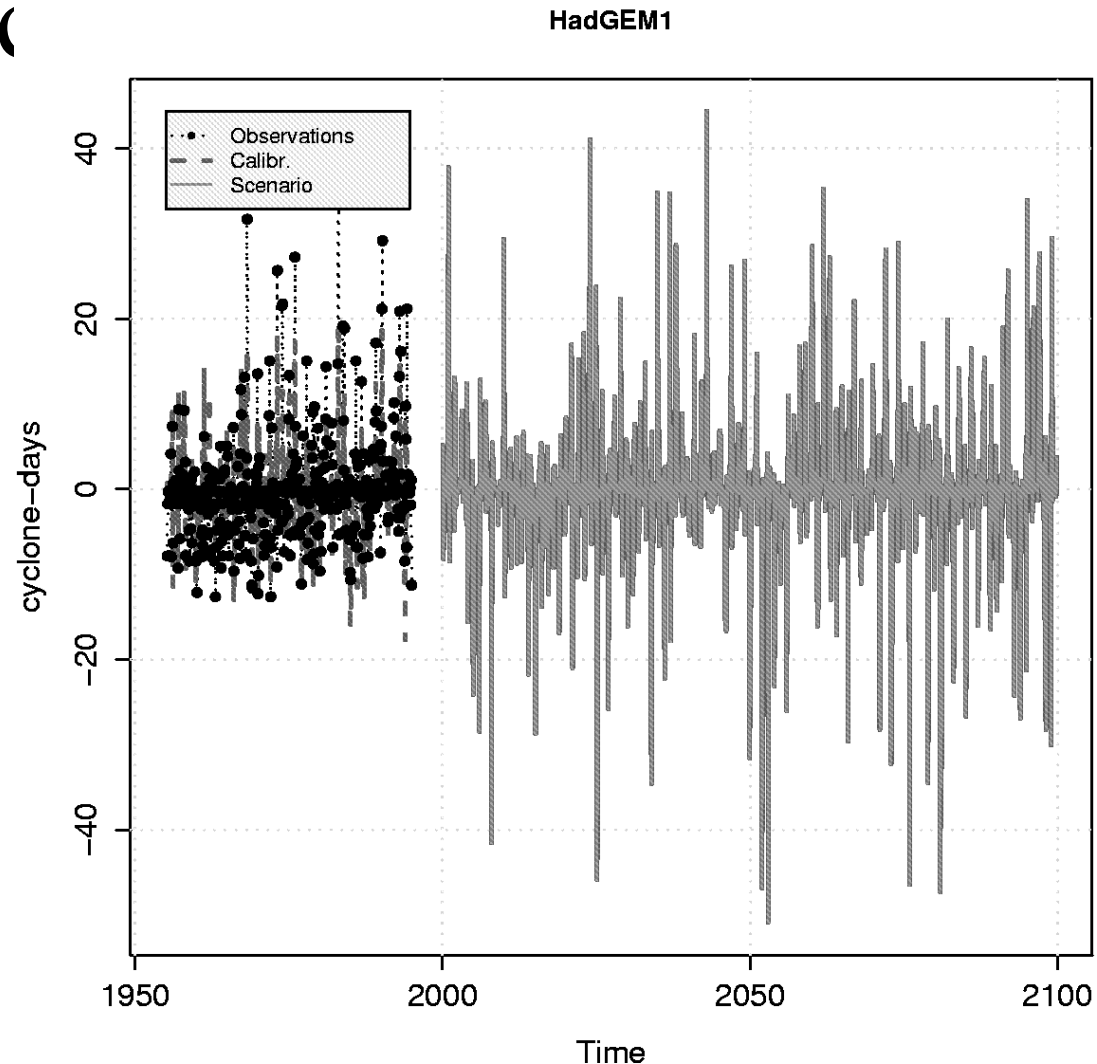


GCMs may not have sufficient spatial resolution for proper representation of cyclones.



Downscaled storm frequency over Fennoscandia

Use the observed time series of cyclone counts as predictand – treating it like a station series – and applying an ordinary downscaling analysis to this, based on monthly SLP.





Summary

- Downscaling as a tool for analysis
 - Methods: strengths and weaknesses
 - Uncertainties
 - Applications
- Extremes
 - Tails of the pdf
 - Record-events & test of non-iid behaviour
 - In the meaning of severe events



Thank you for you attention